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LATE DEVONIAN (FRASNIAN) CORALS FROM CENTRAL DOLPO, NEPAL

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Riassunto. Nel corso di una spedizione geologica italiana nella regione del Dolpo centrale (Nepal), è stata raccolta una piccola fauna a Tabulata e Rugosa, che si è dimostrata di grande interesse sia per la datazione del livello fossilifero (Frasniano Inferiore), sia per le implicazioni paleobiogeografiche. Le calcareniti fossilifere, poste a circa 4900 m di quota, sono ricoperte da un livello ad ooliti e croste ferrose, messo in relazione alla trasgressione del Frasniano Superiore; viene confermata quindi la datazione paleontologica. È stato eretto un nuovo genere, Fruebwirthia gen.n. con specie tipo F. heritschi sp.n.; anche per Scruttonia è stato necessario istituire una nuova specie S. rota sp.n. La fauna comprende anche Kuangxiastraea, Sinodisphyllum, Tabulophyllum (?), Thamnopora, Cladopora, Alveolites e Alveolitella. Dal punto di vista paleobiogeografico questa fauna a Coralli presenta somiglianze sia con quelle dell'Europa occidentale che con quelle della Cina meridionale, fornendo evidenze circa l'esistenza di una estesa provincia faunistica lungo tutto il margine settentrionale del Gondwana. Tale provincia non comprendeva le faune rinvenute nelle rocce coeve del Karakorum, Afghanistan, Iran e Turchia.

Abstract. Rugosa and Tabulata from the Late Devonian of central Dolpo (Nepal) belong to the genera Kuangxiastraea, Scruttonia, Sinodisphyllum, Tabulophyllum (?), Fruehwirthia gen. n., Thamnopora, Cladopora, Alveolites and Alveolitella. The calcarenitic beds yielding the fossils are overlain by a prominent oolitic ironstone with thick ferruginous nodules and crusts. The age of the corals is probably Early Frasnian, thus younger than the Givetian fauna previously reported from eastern Dolpo. Biogeographic relations of the new fauna are with Western Europe and South China.

Introduction (A.T.).

In September 1990, an Italian geological expedition visited the Dolpo region (Himalaya, Nepal) (Garzanti, Nicora & Tintori, 1992), for a project on the Tethys Himalaya concerning the entire Himalayan Range. For many years before 1990 most of that area was closed to westerners; this is the situation again. Geological knowledge

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is very poor, and entirely due to the valuable work by Fuchs (1967, 1977). We could only visit a small area of central Dolpo (Fig. 1), in particular the area surrounding the villages of Do and Tok Khyu (respectively Tarap and Atali in the geological literature). Though small, it is very interesting because it is intermediate between eastern and western Dolpo, which have different sedimentary histories.

The party was mainly interested in the Permo/Mesozoic rocks, but, as much as possible, older rocks were also investigated. A continuous sequence of Devonian to Middle Triassic units crops out at the head of the valley west of Numa La, at about 4900 m. Low tectonic deformation has allowed the fossils to be very well preserved (Nicora & Tintori, 1991; Nicora, 1992). Coral-rich blocks are rather common along the scree, so special attention was paid to the research of the fossils source-rocks. Calcarenite beds overlying a thick dolomite horizon (Garzanti, Nicora & Tintori, 1992) proved to be very rich in organic content, with predominant corals and a few brachiopods. This paper focusses on the corals; they represent a new Himalayan fauna. Only a Givetian fauna was previously known (Fluegel, 1966) from the Devonian of the Tethys Himalaya.



Fig. 1 - Geographic map of the Dolpo area investigated in 1990 by the Italian geological expedition. Location of studied sections is indicated as follows: 1) Tarap (Carboniferous? to Norian); 2) Atali South; 3) Atali-Numa La, yielding the fossils of the present paper; 4) Atali; 5) Atali North; 6) Tarap (Norian to Liassic).

Geological framework and age (A.T.).

The Middle-Late Devonian rocks in Nepal show a wide range of sedimentary environments. From east to west, we can differentiate three major areas: the Manang/Thakkhola/eastern Dolpo subsident terrigenous basin, the western Dolpo and Kumaon dolomitic platform and, in the Spiti-Zanskar Synclinorium, quartzarenitic beaches.

In the first region, the Tilicho Pass Formation consists of pelites with intercalated fine-grained sandstones and carbonates (Garzanti & Pagni Frette, 1991). A Middle Devonian age was inferred from corals and brachiopods for the lower-middle part of this unit in the Thakkhola (Bordet et al., 1967, 1971) and eastern Dolpo (Fluegel, 1966), whereas Frasnian conodonts have been found in its middle-upper part in the Manang area (Fuchs et al., 1988).

In western Dolpo, carbonates (massive dolomites) predominate over calcareous quartzites and quartzarenites (Fuchs, 1977). This unit was called the Muth Formation by Fuchs (1977). In our opinion use of this term is somewhat incorrect, because the Nepalese succession is rather different from the supermature beach quartzarenites found farther west (Muth Quarzite of Spiti and Zanskar; Stoliczka, 1866); these contain very little or no carbonate at all (Gaetani et al., 1986). In any case, the unit yielding our coral fauna cannot be ascribed either to the "Muth Formation" of Fuchs (1977) or to the Tilicho Pass Formation *s.s.*, to which it is, however, broadly related. Furthermore, the transition between the Tilicho Pass Formation and the dolomites of western Dolpo occurs in this area (Fuchs, 1977). Our knowledge of the region is still poor, so we prefer to designate these beds informally as "fossiliferous Late Devonian limestone" (Garzanti, Nicora & Tintori, 1992).

The fossiliferous beds, overlying the Middle ? Devonian dolomites, may well mark the end of a long period in which carbonate sediments were deposited in central and western Dolpo. Our area can be interpreted as the hinge between the terrigenous basin, with centre in the present Thakkhola-Manang region, and the shallow carbonate platform of western Dolpo. In central Dolpo, sedimentation ceased for a while (iron-stone deposition) then started again with shales. This can be related to the late Frasnian transgression in the conodont *gigas* Zone (Talent & Yolkin, 1987). It follows that the fossiliferous limestones are probably pre-late Frasnian.

Kuangxiastraea pengellyi (Milne-Edwards & Haime, 1851) is presently known only from the late Givetian of England; most other species of the genus occur in the Frasnian of Eurasia. Scruttonia and Tabulophyllum (?) implicatum Tsien, 1977 are Frasnian. Sinodisphyllum litvinovitshae (Soshkina, 1949) was described from the late Givetian and early Frasnian of Russia (Ural) and S. China. The new genus Fruehwirthia and the tabulate corals have a less biostratigraphic value. The fauna, as a whole, is most likely of early Frasnian age.



Fig. 2 - Early Late Devonian succession in the Atali-Numa La section. Thick bedded, white dolomites (d) are overlain disconformably by thin-bedded biocalcarenites (b) yielding corals and brachiopods. A condensed ironstone surface (i) follows. Dark pelites (p), of inferred Late Devonian age, hundreds of meters thick, overly the ironstone.

The fact that the two Devonian coral faunas known from Dolpo are of different ages, though from nearby areas, is in accord with great variability in paleoenvironments in the region. Conditions favorable to building of coral reefs probably recurred along the margin of the Thakkhola Basin, the locii varying with time.

Paleobiogeography (A.T. & H.W.F.).

The paleobiogeography of now juxtaposed regions of the Australia-Asian hemisphere (Talent, Gratsianova & Yolkin, 1986; Talent & Yolkin, 1987) was very complex during the Devonian. Several continental blocks with marginal marine facies, can be discriminated by their fossil content (mainly brachiopods), as well as by geological data. Though the Frasnian is normally considered a period of strong cosmopolitanism, Talent & Mawson (1979) pointed out that a measure of provincial heterogeneity can still be discerned. It may represent an inheritance from the marked provincialism of the Early-early Middle Devonian (Ferrari, 1968; Vai, 1991). After the late Frasnian transgression, more cosmopolitan faunas occurred, following the generalized drowning of reefs and carbonate platform (Ferrari & Vai, 1973; Vai, 1991). Faunal provincialism during early Frasnian is consistent with the existence of marine waters between the various blocks along the northern edge of Gondwana.

As already noted above, the north-west Himalayan Devonian, consisting almost entirely of beach deposits (Muth Formation) rarely yields marine faunas. Normal marine sediments are found mainly in the eastern Himalaya, from where the two small coral faunas under discussion are known.

Devonian corals are present in Chitral (Pakistan) in a well-known sequence described by Hayden (1915) and Desio (1966). According to Talent & Mawson (1979), the corals described by Schouppé (1965) from Kuragh Spur and Shogram must be ascribed to the Frasnian rather than to the Givetian. Furthermore, the fossiliferous beds lie a few meters below a thick (5 m) oolitic ironstone, in a position similar to that of our new material from Dolpo. Since the ironstones are probably due to global changes (Garzanti, 1993), the two faunas are likely coeval. Comparison of the faunas shows no significant similarities; a few species from Chitral, however, also occur in European faunas, i.e. Grypophyllum (Leptoinophyllum) vermiculare (Goldfuss, 1826), Macgeea (Macgeea) multizonata (Reed, 1922), Thamnopora reticulata (de Blainville, 1830) and T. boloniensis (Gosselet, 1877) (see Schouppé, 1965). A few Chitral species are also widespread in Russia and Siberia, i.e. Macgeea (Macgeea) multizonata, Thamnopora reticulata and T. boloniensis. According to Desio (1966) the brachiopods, found together with corals in Chitral, are similar to those of West Afghanistan. Our Dolpo fauna does not resemble the coeval faunas from Afghanistan, nor those from northern Iran and southern Turkey (Brice, 1970; Fluegel, 1955; Fluegel & Fluegel, 1961; Gods, 1982; Hubmann, in press). During most of the Paleozoic, continental blocks, such as some of those of Tibet, the Karakorum, Afghanistan, Iran and Turkey, are thought to have been separate from the northern margin of Gondwana, though rather close to it (Leith, 1982; Scotese & McKerrow, 1990). Nonetheless, both brachiopods and corals are strongly provincial, apparently pointing to seas of appreciable width in between them. Also paleoecological factors could be important in separing bioprovinces, especially when long and narrow seas cut landmasses. Paleogeography of the Asian Devonian is still far from being definitively understood. For example, the late Eifelian (Middle Devonian) brachiopod fauna from Burma is astonishing similar to western European faunas raising questions about the relative position of those regions at that time (Talent, Gratsianova & Yolkin, 1986).

A tentative restoration of Frasnian coral biogeography suggests a province extending along the northern edge of Gondwana from South China to west Europe through Burma, north-western Australia, Tethys Himalaya and North Africa. Other provinces are assumed to be related to blocks isolated from the main Gondwana landmass. The Middle Devonian coral fauna from central Dolpo (Fluegel, 1966), is also characterized by western European forms, e.g. *Cyathophyllum dianthus* Goldfuss, 1826, *Neostringophyllum* aff. *N. concavum* Walther, 1928, *Stringophyllum isactisor* (Frech, 1886) and *Favosites goldfussi* d'Orbigny, 1850.

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Systematic paleontology (H.W.F.)*

Subclass **Rugosa** Milne-Edwards & Haime Family *D i s p h y l l i d a e* Hill, 1939 Subfamily *H e x a g o n a r i i n a e* Bulvanker, 1958 Genus *Kuanxiastraea* Yu & Kuang, 1983 Type species *Haplothecia (Kuanxiastraea) elegans* Yu & Kuang, 1983.

1983 Haplothecia (Kuanxiastraea) Yu & Kuang, p. 255.

Diagnosis. See Pedder (1986, p. 655).

Remarks. Kuanxiastraea was erected by Yu & Kuang as a subgenus of Haplothecia. Pedder (1986) regarded Kuanxiastraea as a genus of its own and included it within the subfamily Hexagonariinae Bulvanker. McLean (1989) pointed to the similarities between Kuanxiastraea and Billingsastraea (Sichuanastraea) (He, 1978). They thought that both genera could be synonyms. Against that idea is the morphology of the septa, which are peripherally reticulate or degenerate in Kuanxiastraea, and smooth and complete in Sichuanastraea. Species of Kuanxiastraea have been reported from Middle and Late Devonian strata of China, Australia, the Rudny Altai Region and Western Europe. According to Kursten (1987), the unique species of Kuanxiastraea of northern America, K. julli Pedder, 1986, belongs to Scruttonia.

Kuanxiastraea pengellyi (Milne-Edwards & Haime, 1851)

Pl. 1, fig. 1, 2

1851 Smithia pengellyi n. sp. Milne-Edwards & Haime, p. 422. 1967 Haplothecia pengellyi - Scrutton, p. 274, pl. 42, fig. 1-5; pl. 43, fig. 1,2 (cum syn.). 1986 Kuanxiastraea pengellyi - Pedder, p. 655.

Material. Specimens n.6760 (ex AD30/1, 16 field numbers) and 6761 (ex AD30/5, 7 field numbers).

Description. Colonies thamnasterioid, locally aphroid, with maximum diameter 90×70 mm and thickness 20 mm; centres of adjacent corallites 7 to 18 mm apart; diameter of tabularia 3 to 5 mm; diameter of corallites 7 to 19 mm. Larger corallites have 18 to 21 x 2 septa, smaller corallites 12 to 15 x 2 septa. Major septa extend to the axis and could be connected. Minor septa thin and confined to the dissepimentaria. Marginal septa degenerate to a network or end in an aphroid dissepimentarium. Both orders of septa are equipped with yardarm carinae. Near the inner edge of the dissepimentaria, septa are dilated. Tabulae are complete to incomplete. Near the dissepiment

^{*} The systematic arrangement of higher taxa follows Hill (1981).

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tarium/tabularium boundary dissepiments are smaller than in other parts of the dissepimentarium. Horseshoe dissepiments lacking.

Remarks. The specimens are very similar to K. pengellyi described by Scrutton (1967) from the Givetian of England.

Genus Scruttonia Cherepnina, 1974

Type species Smithia bowerbanki Milne-Edwards & Haime, 1851.

1974 Scruttonia gen.n. Cherepnina, p. 202.

1978 Billingsastraea (Sichuanastraea) n.subgen. He, p. 134.

1987 Scruttonia Kuester, p. 36 (cum syn.).

Diagnosis. See Hill (1981, p. F284), Rohart (1982, p. 105), McLean (1989, p. 243).

Discussion. Cherepnina (1974) erected the genus for colonial corals discriminated from Frechastraea Scrutton only by having a thamnasterioid corallum. The taxonomic status of Scruttonia is debatable. Roemer (1883) and Frech (1885) included the type species in Phillipsastrea d'Orbigny, Soshkina (1951, 1952) in Pachyphyllum Milne Edwards & Haime, Pickett (1967) and Coen-Aubert (1974) in Billingsastraea Grabau and Scrutton (1968) and Birenheide (1978) in Frechastraea. Birenheide (1990) regarded Scruttonia as a subgenus of Frechastraea, whereas Coen-Aubert (1980), Rohart (1982), Cotton (1983), Pedder (1986) and McLean (1989) agreed with the conception of Scruttonia as a discrete genus, as Cherepnina (1974) had done. In this connection the question of the presence or absence of horseshoe dissepiments in the type species is pivotal. The original of S. bowerbanki is apparently lost (Scrutton, 1968, p. 254). Description and illustration of the type throws no light on this problem. Real horseshoe dissepiments are structurally very complicated. The dissepiments are sometimes not only bent, horseshoe-like, in the radial direction of the septa, but also in the normal direction. Placed one upon the other, they make up a ring- and wall-like structure within or peripheral to the dissepimentarium. They are morphogenetically older than the adjacent normal dissepiments and/or tabulae. Genera such as Macgeea, Thamnophyllum, and Trapezophyllum, of the family Phillipsastreidae are good examples of this type (Schouppé, 1958). Half or false horseshoe dissepiments also occur occasionally in genera of Disphyllidae. True horseshoe-ring-structures at Scruttonia fail, but casual false horseshoe-dissepiments exist.

Scruttonia bowerbanki (Milne-Edwards & Haime, 1851)

Pl. 2, fig. 1, 2

1851 Smithia bowerbanki n. sp. Milne-Edwards & Haime, p. 423. 1980 Scruttonia bowerbanki - Coen-Aubert, p. 5, pl. 1, fig. 1-3 (cum syn.). 1986 Scruttonia bowerbanki - Pedder, p. 658.

Material. Specimens n. 6754 (ex AD28/9) and 6755 (ex AD28/11).

Description. Corallum thamnasterioid, disc-like, with a diameter of more than 100 mm, and a thickness of more than 30 mm. Diameters of the circular tabularia are 1.0 to 1.5 mm, with distances centre-to-centre between 3 and 5 mm. Number of tabularia in 100 mm² varies between 6 and 8. There are 10×2 septa. Within the tabularia, major septa are thin and often reach the centre of the tabularia where they are connected. Minor septa are well developed, crossing only the dissepimentaria. Both septal orders are spindle-like, dilated at the dissepimentarium/tabularium boundary, but only slightly dilated and mostly bent within the dissepimentaria. Yardarm carinae are well developed. Dissepimentaria regular, with mostly small dissepiments; horseshoe dissepiments are absent. Seven to eight tabulae occur per mm; they are horizontal or dip slightly upwards towards the axis. Fans of trabeculae are present near the dissepimentarium/tabularium boundary.

Remarks. The specimens display characteristic alteration of light and dark growth strips; these are visible in both transverse and longitudinal sections. The longitudinal section shows that these strips are related with small zones of dissepiments. *S. bowerbanki* is known from the Frasnian of western Europe.

Scruttonia rota sp. n.

Pl. 2, fig. 3, 4

Holotype. N. 6749 (ex AD28/1).

Locus typicus. Atali-Numa La section, central Dolpo, Nepal.

Stratum typicum. "Fossiliferous Late Devonian Limestone" of Garzanti, Nicora & Tintori (1992), early Frasnian.

Paratype. Specimens n. 6750 (ex AD28/4) and 6753 (ex AD30/2).

Derivatio nominis. rota (lat.) - wheel, referring to the wheel-like appearance of the tabularia in transverse section.

Diagnosis. A species of *Scruttonia* with central hub-like linkage of the major septa, with diameter of tabularia to 2 mm, and with 12×2 septa.

Description. The massive, flattened, thamnasterioid colonies are up to 100 mm in diameter and 30 mm in thickness. Diameters of the tabularia 1.8 to 2.0 mm. In 100 mm² there are 5 to 8, mostly 6 tabularia. There are 10 to 12 x 2 septa. Major septa are long and thin, connected axially and, together with the stereozone, forming a wheel-like structure around the tabularium. The hub-like connecting point of the septa is seen in longitudinal section as a pillar within the tabularia. Both order of septa are spindle-like, dilated around the tabularia and, partly, wall-like, connected to a stereozone. The minor septa do not project into the tabularia. Within the dissepimentaria the septa are slightly thicked and variably carinate. The tabulae are complete, flat and average six per mm. The dissepimentaria consist of small, globose dissepiments. Horse-shoe-dissepiments are lacking. Trabeculae are arranged in asymmetrical fans.

Remarks. The new species is characterized by the wheel-like appearance of the tabularia in transverse sections. The measurements are similar only to *S. sanctacrucensis* Rozkowska, 1953 and *S. bolonensis minor* Coen-Aubert, 1980. Both species do not, however, show the characteristic wheel-like feature.

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Family Cyathophyllidae Dana, 1846 Subfamily Zaphrentinae Milne-Edwards & Haime, 1850

Genus Sinodisphyllum Sun, 1958

Type species Disphyllum (Sinodisphyllum) variabile Sun, 1958.

1958 Disphyllum (Sinodisphyllum) gen n. Sun, p. 11.

? 1961 Mansuyphyllum gen. n. Fontaine, p.100.

? 1975 Aristophyllum n.g. Bulvanker, Spasky & Kravtsov in Ivanovsky, p. 78.

Diagnosis. See Liao & Birenheide, 1989, p. 88.

Sinodisphyllum litvinovitshae (Soshkina, 1949)

Pl. 4, fig. 1

1949 Campophyllum litvinovitshae sp. n. Soshkina, p. 86, pl. 38-40, fig. 1-3. 1977 Sinodisphyllum litvinovitshae - Liao, p. 43, pl. 1, fig. 6-8; pl. 3, fig. 3 (cum syn.). 1989 Sinodisphyllum litvinovitshae - Liao & Birenheide, p. 88, pl. 1, fig. 1-6.

Material. Specimens n. 6757 (ex AD28/8) and 6758 (ex AD30/3).

Description. The diameter of the ceratoid corallites is 16 to 22 mm. The 28 radially arranged major septa are 4.0 to 5.5 mm long. The cardinal septum is a little withdrawn and shorter than adjacent septa. The diameter of the free axial space is 3 to 5 mm. Minor septa are about half as long as the major septa, and terminate inside the tabularium. Both orders of septa are wedge-like within the dissepimentarium and thin out in the tabularium. Minor septa, within the dissepimentarium, are a little thinner than the major septa. Eight to 10 rows of small dissepiments, occasionally of herring-bone-type, occur.

Remarks. The long, wedge-shaped septa, the shorter cardinal septum, the length of the minor septa, and the numerous rows of dissepiments are differences between *Mictophyllum* Lang & Smith, 1939 and *Temnophyllum* Walther, 1928. The corallite diameter and number of septa fall within the range of variation in *S. litvinovitshae*. The species is known from the Givetian and Frasnian of the Ural and South China.

Family Charactophyllidae Pedder, 1972

Genus Fruehwirthia gen. n.

Type species Fruehwirthia heritschi sp. n.

Derivatio nominis. The new genus is named after Josef Fruehwirth, retired laboratory assistant of the Institute of Geology and Paleontology of the University of Graz. Without his valuable help through many years my scientific work in paleontology would have been impossible.

Diagnosis. Corallum ceratoid with two septal orders peripherally, with intraseptal and interseptal dissepiments; tabulae incomplete and axially arched.

Remarks. The new genus recalls somewhat *Ceciliaphyllum* McLean, 1982 from the Frasnian of British Columbia; differences are the inverted dissepimentarium and the existence of degenerate septa in that genus. On the other hand the new genus may be compared with *Federowskicyathus* Rozkowska, 1980, especially *F. radicensis* McLean & Pedder, 1984 from the Frasnian of Western Canada. This species is characterised by intra and interseptal dissepiments. Differences are the existence of well developed presepiments and the tendency of the septa to be twisted axially in *Federowskicyathus*. Because of the occurrence of intraseptal dissepiments, the systematic position of the new genus is obscure.

Fruehwirthia heritschi gen. n., sp. n.

Pl. 3, fig. 1-4; Pl. 4, fig. 2

Holotype. Specimen n. 6751 (ex AD28/6).

Locus typicus. Atali-Numa La Section, central Dolpo, Nepal.

Stratum typicum. "Fossiliferous Late Devonian Limestone" of Garzanti, Nicora & Tintori (1992), early Frasnian.

Paratype. Specimen n. 6752 (ex AD28/10).

Derivatio nominis. I dedicate the new species to my late teacher Univ. Prof. Dr. Franz Heritsch, who inspired my love for Palaeozoic corals.

Diagnosis. Diameter of the solitary corallite below the base of the calice around 20 mm, with 27-29 x 2 septa. Major septa extend into the tabularium; minor septa are half as long as major septa. Both orders of septa have intraseptate dissepiments.

Description. Ceratoid coralla about 60 mm long, with diameter 18 to 23 mm. The transverse section shows 27 to 29 x 2 septa. The development of the major septa is different: in the holotype, one septum is long and connected in the axial region with sections of tabula; other major septa are distinctly shorter. The other specimen (paratype) shows long major septa, twisted or bent in the axial region. The minor septa are half or less than half as long as the major septa. In the peripheral part of the corallite, both orders of septa are wedge-like, dilated and fork-like, split into two branches. As well as V-shaped interseptal dissepiments ("Hohlkehlen-Dissepiment" of Birenheide, 1963) there are, within these forks, intraseptal dissepiments. At the boundary of the dissepimentarium/tabularium, the dissepimentarium, up to 3 mm wide, consists of small, oblique, elongated dissepiments. The tabularium is formed of an axial part of table-shaped, partly incomplete tabulae and a peripheral part of great, cyst-like tabellae.

Remarks. The two specimens differ only in development of the major septa, but both come from the same locality, have the same number of septa and the same diameter.

Genus Tabulophyllum Fenton & Fenton, 1924

Type species Tabulophyllum rectum Fenton & Fenton, 1924.

1924 Tabulophyllum Fenton & Fenton, p. 30. 1928 Apolythophyllum Walther, p. 35.

Diagnosis. See Sorauf, 1987, p. 678.

Remarks. The characteristics of the genus are elongate presepiments, long major and short minor septa and an axial region without septa. These morphological characteristics vary considerably (Sorauf, 1988).

Tabulophyllum (?) implicatum Tsien, 1977 Pl. 4, fig. 3, 4

1977 Tabulophyllum implicatum sp. n. Tsien, p. 271, text-fig. 8, pl. 4, fig. 3-5.

Material. Specimen n. 6759 (ex AD28/2).

Description. The corallite is ceratoid and more than 4 cm long. At a diameter of 22 mm are 34 x 2 radially arranged septa. Major septa are long, dilated toward the periphery, thin and twisted axially, and form, with slightly domed tabulae, a loose axial structure. Cardinal septum (?) is short. Minor septa are 2 to 3 mm long. The peripheral parts of the septa are connected forming a 1.0 to 1.5 mm thick wall. The dissepimentarium is narrow and consists of one or two rows of inclined, elongate dissepiments. Presepiments lacking. Tabulae are flattened, complete or incomplete, and occasionally slightly domed axially.

Remarks. The characteristics of the specimens are the long major septa swirling around the axis and generating with the tabulae a loose axial structure. Short cardinal septum (?). Dissepimentarium narrow with few rows of dissepiments. Our specimens differ from the holotype of *T. implicatum* in the lack of a stereozone. The occurrence of such a zone of a stereoplasmatic coating of dissepiments is assumed to depend on growth conditions and thus primarily on ecological factors.

The generic identification of the species is uncertain (Birenheide, 1978, p. 66). Lack of presepiments, the existence of a weak axial structure and a short cardinal septum indicate clearly that "T. implicatum" does not belong within the genus Tabulo-phyllum.

Subclass **Tabulata** Milne-Edwards & Haime, 1850 Order Favositida Wedekind, 1937 Suborder Favositina Wedekind, 1937 Family *Pachyporidae* Gerth, 1921 Subfamily *Pachyporinae* Gerth, 1921 Genus *Thamnopora* Steininger, 1831 Type species *Alveolites cervicornis* de Blainville, 1830.

1831 Thamnopora gen. n. Steininger, p. 10

Remarks. The morphological characteristics of thamnoporoid *Tabulata* are the diameters of the branches, corallites and pores, the spacing of mural pores and tabulae, the thickness of walls, and the presence of spines. Most of these characteristics are

highly variable and their measurements depend on orientation of the section. These uncertainties are probably one of the causes why, according to Birenheide (1985) 70 to 80 species, and according to Nowinski (1976) and myself, more than 150 species of the genus have been described from the Silurian and Devonian. This great number of species makes unequivocal determination difficult, and revision of the genus necessary.

Thamnopora cf. urensis urensis Dubatolov, 1959

Text-fig. 3

cf. 1959 Thamnopora urensis Dubatolov, p. 96, pl. 30, fig. 3-6.

Material. Specimens n. 6748 (ex AD30/4) and 6764 (ex AD28/5,7,14,17,18).

Description. The colonies are fascicular, with cylindrical branches of diameter 7.5 to 11 mm. In transverse section the corallites are polygonal, with round or suboval lumen. The diameter of the lumen is axially 0.3 to 0.7 mm, peripherally 1.5 to 1.7 mm. In longitudinal sections the corallites of the axial zone are vertical and, in the peripheral zone, arcuate and oblique to the growth axes of the branch. The micro-



Fig. 3 - Thamnopora cf. urensis Dubatolov, 1959. Specimen n. 6748 (ex AD30/4) from Atali-Numa La section (central Dolpo, Nepal); x 7.

Late Devonian corals, Nepal

structure of the walls is fibro-lamellar. The dark median suture is distinct. The thickness of the walls is 0.1 to 0.3 mm axially, 0.4 to 1.0 mm peripherally. The calices are deep and funnel-like. Tabulae are rare, thin and horizontal. The connecting pores are round, arranged in a single row, and of diameter 0.1 to 0.2 mm, spaced 0.4 to 1.0 mm apart. Septal spines are fairly developed. In longitudinal sections the spines could be sometimes to 0.7 mm long and squamulae-like. In deeper parts of the calices, the spines are sometimes connected, forming "toothed" septal lamellae.

Remarks. The existence of septal lamellae in deeper parts of the calices is not characteristic for *Thamnopora*. They recall *Parastriatopora* Sokolov, but this genus is characterised by hexagonal axial corallites and pronounced thickness of the peripheral regions of the walls.

Suborder Alveolitina Sokolov, 1950 Family *A l v e o l i t i d a e* Dunkan, 1872 Genus *Alveolitella* Sokolov, 1952 Type species *Alveolites fecundus* Lecompte, 1939.

Diagnosis. See Sokolov, 1952, p. 60.

Remarks. Sokolov (1952) erected Alveolitella for ramose Tabulata with polygonal, thin-walled corallites in the axial zone and semi-lunar to round-triangular, thick-walled corallites towards the periphery. Iven (1980), Hill (1981), Byra (1983) and Birenheide (1985) united Alveolitella with Alveolites Lamarck. Differences between both genera are the form of the colonies, the form of the corallites and thickness of the walls. The genus is known from the Devonian of Europe, Russia and China, with about 20 species.

> Alveolitella cf. rarispinosa Nowinski, 1976 Pl. 1, fig. 3

1976 Alveolitella rarispinosa sp.n. Nowinski, p. 63, pl. 8, fig. 2-5.

Material. Specimens n. 6762 (ex AD28/3,12) and 6763 (ex AD28/13).

Description. Colonies are ramose, up to 90 mm long, with a diameter of 8×15 mm. Corallites of the axial zone are polygonal to subpolygonal, with maximum diameter 0.4 mm. Towards the periphery they are slightly bent, becoming gradually alveolitid, and opening obliquely to the surface with diameter of 0.25 x 0.6 mm. Wall thickness in the axial zone is 0.1 mm and in the peripheral zone 0.3 mm. The dark median suture is distinct and semi-lunar. Connecting pores are large and round. Septal spines are rare. The tabulae are thin and slight bent.

Remarks. The specimens are similar to *A. rarispinosa* from the Frasnian of Poland. They differ from this species in the form of the septal spines, and in thickness of the wall at the periphery.

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PLATE 1

- Fig. 1 Kuanxiastraea pengellyi (Milne-Edwards & Haime, 1851). Transverse section. Specimen n. 6761/7; x 3.25.
- Fig. 2 Kuanxiastraea pengellyi (Milne-Edwards & Haime, 1851) Longitudinal section. Specimen n. 6761/5; x 2.92.
- Fig. 3 Alveolitella cf. rarispinosa Nowinski, 1976. Transverse section. Specimen n. 6763; x 3.25.

All specimens from Atali-Numa La section (central Dolpo, Nepal).

PLATE 2

- Fig. 1 Scruttonia bowerbanki (Milne-Edwards & Haime, 1851). Transverse section. Specimen n. 6754; x 2.92.
- Fig. 2 Scruttonia bowerbanki (Milne-Edwards & Haime, 1851). Longitudinal section. Specimen n. 6754; x 2.92.
- Fig. 3 Scruttonia rota sp.n. Holotype. Transverse section. Specimen n. 6749; x 3.25.
- Fig. 4 Scruttonia rota sp. n. Longitudinal section. Specimen n. 6750; x 5.85.

All specimens from Atali-Numa La section (central Dolpo, Nepal).

PLATE 3

Fruehwirthia heritschi gen. n, sp. n.

Fig. 1 - Holotype, specimen n.	6751.	Transverse sect	ion; x 3.90.
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- Fig. 2 Holotype, specimen n. 6751. Transverse section; x 9.10.
 Fig. 3 Holotype, specimen n. 6751. Transverse section; x 22.75.
- Fig. 4 Holotype, specimen n. 6751. Longitudinal section; x 3.90.

All specimens from Atali-Numa La section (central Dolpo, Nepal).

PLATE 4

- Fig. 1 Sinodisphyllum litvinovitshae (Soshkina, 1949). Transverse section. Specimen n. 6757; x 4.55.
- Fig. 2 Fruehwirthia heritschi gen. n., sp. n. Transverse section. Specimen n. 6752; x 4.55.
- Fig. 3 Tabulophyllum(?) implicatum Tsien, 1977. Longitudinal section. Specimen n. 6759; x 4.55.
- Fig. 4 Tabulophyllum(?) implicatum Tsien, 1977. Transverse section. Specimen n. 6759; x 3.57.

All specimens from Atali-Numa La section (central Dolpo, Nepal).

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