

NEW INSIGHTS ON *ANTHRACOTHERIUM MONSVIALENSE* DE ZIGNO, 1888 (MAMMALIA, CETARTIODACTYLA) FROM THE LOWER OLIGOCENE OF MONTEVIALE (VICENZA, NORTHEASTERN ITALY)

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Abstract. In Italy, anthracotheres are represented by a few fossils, most of them described during the XIX century and without a standardized scientific method. Anthracotherium monsvialense De Zigno, 1888 was originally erected from a fossil discovered in the site of Monteviale (Vicenza, northeastern Italy), whose Rupelian (MP21) lignitic beds yielded the richest lower Oligocene evidence of the genus Anthracotherium in Europe. A. monsvialense ranges from MP21 to MP23 and its small size has been interpreted as a consequence of the insular environment, at least at Monteviale.

In this study, we summarize the long history of Italian findings providing new descriptions of dental and postcranial morphological features of *A. monsvialense*, and comparing such small anthracothere with its Asian and European relatives. Morphometric analyses are also performed on teeth, in order to verify the presence of evolutionary trends of the genus *Anthracotherium*.

INTRODUCTION

The origin and evolutionary history of the Cetartiodactyla (Montgelard et al. 1997) is far to be completely solved. They appeared across the Holartic region at the beginning of the Eocene (e.g., Rose 1996; Blondel 2001), from uncertain centers of origin, during the Mammalian Dispersal Event (MDE), one of the key events of the Paleocene/ Eocene boundary (Aubry et al. 2007). The MDE introduced the earliest members of the orders Cetartiodactyla, Perissodactyla and Primate (APP taxa) in North America, Asia and Europe during a brief warming event (100-200 kyr) which occurred in the basal Eocene, the "Paleocene/Eocene thermal maximum" (Koch et al. 1992; Bowen et al. 2001; Gingerich 2006). According to recent stratigraphic calibrations, the APP taxa appeared earlier in NW Europe (MDE phase I), in the latest Paleocene, suggesting an older evolution of the groups, possibly migrated from mid-latitude Asia (Hooker 2015). It implies that the presence of Diacodexis gigasei, Smith, Smith and Sudre, 1996 within the latest Paleocene PE I zone in NW Europe represents so far the oldest record of order Cetartiodactyla (Hooker 2015),

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being previously unknown before the Eocene (Prothero & Foss 2007).

Among Cetartiodactyla, the family Anthracotheriidae includes even-toed and non-ruminant herbivores, with five cusps on upper brachyodont molars, a lophed structure which prefigures the selenodonty adapted to a frugivorous/folivorous diet (more evident on the higher-crowned lower molars), and separated cuboid and navicular among the tarsals (Janis 1995; Blondel 2001).

Most of the recent interest about anthracotheres concerns the relationships of this taxon with hippos and cetaceans (Montgelard et al. 1997; Agnasson & May-Collado 2008). In fact, while Pickford (2008) rejected the hypothesis in favor of a divergence from the suid family of family Palaeochoeridae, cladistic analyses confirm the affinities between Hyppopotamoidea and anthracotheres, proving the speciation of the former taxon probably from bothriodontines (Boisserie et al. 2011; Lihoreau et al. 2015).

The family Anthracotheriidae is represented by a huge number of well-differentiated species since their appearance in the uppermost middle Eocene Asian records (Pondaung Formation, Tsubamoto et al. 2002; Tsubamoto & Tsogtbaatar 2008; Soe 2008; Ducrocq et al. 2015). Members of the family have been reported in Asia (Ozansoy 1962; Tsubamoto & Tsogtbaatar 2008), Europe (Ozansoy 1962; Ducrocq 1995; Legendre 1995; Antoine et al. 2011; Főzy & Szente 2014), Africa (Ducrocq et al. 2001; Miller et al. 2007; Holroyd et al. 2010; Sileem & Hewaidy 2015), North and Central America (Kron & Manning 1998; Rincon et al. 2013), spreading in Europe since the late Eocene with the genera Elomeryx, Doplopus and Prominatherium (Kowalevsky 1893; Hellmund 1991; Meinolf & Kurt 1994; Kostopoulos et al. 2012; Lihoreau et al. 2009). The single K/Ar dating of 37.7 ± 1.5 Ma (early Priabonian, late Eocene, MP16) for the anthracotheres from Dětăn, in Czech Republic (Fejfar 1987; Fejfar & Kaiser 2005) has been revised by Aguilar et al. (1997), who assigned the fauna to the MP21, the isotopic da-ting being possibly falsified by argon excess (Kelley 2002).

The genus *Anthracotherium* appeared in Europe in correspondence of the "*Grande Coupure*" faunal turnover (Mammalian Paleogene Zone - MP21) (Stehlin 1910a). In fact, the climatic changes and the formation of land connections at the Eocene-Oligocene boundary (Stehlin 1910b; Lihoreau et al. 2004) through a south-east Balkans-Anatolian way (Ducrocq 1995) facilitated the migration and probably the speciation and divergence of new species of anthracotherids. At the end of Oligocene, the genus completely disappeared in Europe (Lihoreau & Ducrocq 2007; Scherler et al. 2010) with the latest occurrence recorded at Rickenbach (MP29, Switzerland) (Becker et al. 2004; Mennecart et al. 2012).

This work is focused on the richest lower Oligocene findings of the genus Anthracotherium in Europe, discovered during mining activity in the surroundings of Monteviale village (Vicenza, northeastern Italy). The species Anthracotherium monsvialense De Zigno, 1888 had been reported in Italy, dubitatively at Cadibona and in few localities in southern France, western Switzerland, Germany, Spain and Turkey (Lebkuchner 1974, tab. 23-1; Kotsakis 1986; Sudre 1995; Becker et al. 2004; Scherler 2011, 2013) and spans from early to early "middle" Oligocene (Sudre 1995; Pandolfi et al. 2016). Since about MP25, Anthracotherium magnum Cuvier, 1822 occurred in Europe, becoming quite frequent in the European Oligocene record (Lihoreau & Ducrocq 2007).

Our detailed description of the fossils of *A*. *monsvialense* from the type locality provides specific

morphological information of teeth and postcranial anatomy, improving the knowledge about this taxon within the family Anthracotheriidae. In fact, the detailed knowledge of dental and postcranial characteristics of the species is of primary importance in order to enrich the extant knowledge about anthracotheres dispersal pathways in Europe (Lihoreau et al. 2004). The occurrence of *A. monsvialense* in the Venetian region at the *Grande Coupure* suggests a land connection via an island chain, consisting of micro-continents and volcanic back-arcs between southern Europe and southeastern Asia, where the genus *Anthracotherium* probably originated (Tsubamoto et al. 2002; Böhme et al. 2014; Pandolfi et al. 2016).

AN OVERVIEW ON ITALIAN ANTHRACOTHERES

The first record of Italian anthracotheres was reported by Borson (1820), who described a few teeth of a completely unknown animal at the Regia Accademia delle Scienze di Torino. Such teeth were found in the lignite deposits near the village of Cadibona (Savona, northwestern Italy), which yielded the oldest and richest collection of *A. magnum* in Italy (Borson 1820; Squinabol 1890a, 1890b; Sieber 1935). Two years after Borson's report, Georges Cuvier erected the genus *Anthracotherium* based on the fossils from Cadibona (Cuvier 1822). The name, formed by the Greek words $\ddot{\alpha}\nu\vartheta\rho\alpha\xi$ = 'coal' and $\vartheta\eta\rho$ íov = 'beast' (Agassiz 1842), was clearly inspired by the sedimentary context where the fossil was found.

In addition to Cadibona's specimens, a few remains had been reported from Agnana (Reggio Calabria) by Montagna (1857), and determined as *A. magnum* by Flores (1897). Even though the ectostyle is well developed on M3, suggesting a correct determination of the few teeth remains from Agnana (Flores 1897; plate 1 fig. 5), the material is in need of a careful systematic revision and standardized description. The age of these fossils was estimated between late Eocene and early Miocene (Dal Piaz 1929; Esu & Kotsakis 1983; Kotsakis 1986).

Since 1858, several remains of anthracotheres have been found in the eastern side of Alps, in the Vicenza province, at Zovencedo and Monteviale (De Zigno 1888), and, less than one century later, at Chiuppano (Leonardi 1950; Accordi 1951).



Fig. 1 - A) location map of Monteviale with the geologic detail of the Berici Hills (white bricks indicate the area of Castelgomberto Limestone, black zones are urban centers); B) reconstruction of depositional context during the Oligocene (schemas from Frost 1981, modified).

The fossils from Zovencedo were found in a lignite bed with "tortoises and other animals" and had been described by Francesco Secondo Beggiato (1806-1883) in 1865. Unfortunately, in 1945, an Allied bombing partially destroyed the historical building of the Museo Naturalistico Archeologico of Vicenza, where the fossils were housed, and the material today does not result in the inventory database of the Museum, indicating that the Zovencedo collection is completely lost (Dal Lago, pers comm., 2015). Recently, Bona and Grandi (2014) described a new maxillar fragment of Anthracotheridae (Prominatherium cf. dalmatinum) from Grancona (Vicenza), suggesting a late Eocene age for the fossil. Finally, the remains of anthracotheres reported by Stehlin (1910a) and Kotsakis (1986) from Tuscany and from Oschiri (Sardinia) at the MN2 (lower Miocene) should be considered with caution (Bruijn & Rümke 1974; Kotsakis 1986; Made 1999, 2008).

THE MONTEVIALE SITE: GEOLOGICAL AND PALEONTOLOGICAL CONTEXT

Geological context. In the Veneto region (northeastern Italy), Oligocene sedimentary rocks are well represented along the eastern Lessini Mountains margin and in the Berici Hills (Rasser & Harzhauser 2008). During the early Oligocene, the area was characterized by a shallow-water carbonate platform, delimited northward by emerged land and southward by a coral reef separating a welloxygenated lagoon from a deep basin (e.g., Frost 1981) (Fig. 1). This sedimentary environment is represented by the richly fossiliferous Castelgomberto Calcarenites, widely outcropping in the study area (e.g. Ungaro 1978; Frost 1981; Mietto 1988). The lagoonal sedimentation of Castelgomberto Platform was locally interrupted by explosive volcanism, which often lasted for only one paroxistic episode (Mietto 2006). Some explosive volcanic structures emerged from the shallow sea (ca. 20 m deep) forming short-lived islands. Some of these islands, like Zovencedo (Berici Hills) and Monteviale (Lessini Mountains), were large enough to sustain a rich vegetation and terrestrial vertebrates (Pandolfi et al. 2016), as testified by their lignite fossiliferous deposits. Such temporary islands were subsequently submerged and covered by new carbonate deposits that eventually filled the lagoon (Mietto 1988, 2006). After subsequent phases of emersion, volcanism and marine sedimentation, the whole area was uplifted by the Alpine orogeny.

The vertebrate fauna of Monteviale. During the XIX and the first half of XX century, in the area of Monteviale, the exploitation of lignitic beds led to the discovery of a rich Oligocene vertebrate fauna (Mietto 2006 and references therein). The faunal assemblage consists of fishes ("Lepidocottus" papyraceous (Agassiz, 1832)), amphibians (Palaeobatrachus sp.), reptiles (Trionyx italicus von Schauroth, 1865, Bergouniouxchelys vallesnerii (Bergounioux, 1954), Diplocynodon cf. D. ratelii Pomel, 1847), and mammals, represented by Dugongidae Gray, 1821, pantolestid (Epapheliscus italicus Van Valen, 1966 and ?Pantolesta incertae sedis), bats (Archaeopteropus transiens Meschinelli, 1903), and ungulates (Epiaceratherium bolcense Abel, 1910, Anthracotherium monsvialense De Zigno, 1888, Anthracochoerus



Fig. 2 - Achille De Zigno (Padova 1813 - Padova 1892). De Zigno, a famous Italian palaeobotanist, also published fifteen papers about fossil vertebrates of northern Italy and had been member of the Accademia dei Lincei and president of several Italian Institutes (i.e. the Istituto Veneto di Scienze, Lettere ed Arti, the Accademia Galileiana and the Società Geologica Italiana; Roghi et al. 2013). Photo from the family archive of N.H. Alberto Lonigo, Mestrino (Padova).

stehlini Dal Piaz, 1931, Anthracochoerus fabianii Dal Piaz, 1931, and ?*Propalaeochoerus paronae* Dal Piaz, 1930) (Mietto 2006; Pandolfi et al. 2016).

The remains were collected during coal mining activities and dug by miners, who sold the fossils to contemporary naturalists, private collectors and paleontologists (Mietto 2006; Pandolfi et al. 2016). This process prevented the complete recovery of the specimens preserved in the lignite beds. Miners probably sold only the best specimens they came across, and wasted the ones considered not interesting or not recognized as fossils. As the mining activity stopped in 1952, no further discoveries have been made, and the mines' tunnels are now mostly covered. To date, no outcrop of the original lignite beds is known (Mietto 2006).



Fig. 3 - Upper dental series of *A. monstrialense* MGP-PD 26556: (B) is the holotype of the species (A: as drawn by De Zigno 1888), despite the additional last molars and the lower incisors, inserted in the jaw in place of the upper ones (white squares) made by the miner seller in 1886.

The anthracotheres from Monteviale. Anthracotheres from Monteviale had been mentioned for the first time by Beggiato (1865), who described two isolated teeth (a m3 and a lower incisor) and compared them with Anthracotherium from Zovencedo. Beggiato (1865) recognized the presence of a small Anthracotherium morphotype in the site of Monteviale. A palate with the complete dentition of an anthracothere collected at Monteviale was bought by professor Giovanni Omboni (1829-1910) for the Gabinetto of the Regia Università di Padova in 1886 and it had been considered the most important fossil from the site. This fossil represents the holotype of Anthracotherium monsvialense, originally described by Baron Achille De Zigno (Fig. 2) in 1888, and it is currently housed at the Museo di Geologia e Paleontologia, Università di Padova (MGP-PD), inventory number MGP-PD 26556 (Fig. 3).

The holotype of *A. monsvialense* became quickly the object of strong debate (e.g., Dal Piaz 1932) due to the presence of one supernumerary molar on each upper tooth-series (De Zigno 1888; Flores 1897). Twenty years later, Stehlin (1910a) finally recognized the partial forgery of the specimen, consisting in the addition of two supplementary M3s and lower incisors, probably made

Inv. n. upper											
teeth	M3	M2	M1	P4	P3	P2	P1	С	I3	I2	I1
(MGP-PD)											
27361 left	0	0									
27361 right	0	0									
31495	1	2									
31489				2	2+						
31479					2+	2+					
27363	1	2	4	3-	1						
31486	1	3-	3								
27359 left	1+	4									
31501	2	3									
31413	3	3+									
27364 right	3	4-									
31502		4	5								
27364 left	3	4/5	5								
27366	5	4	2								
27360 left								2/3	3/4	2/3	
27360 right										2/3	2
31503									3	3	
Inv. n. lower											
teeth	m3	m2	m1	p4	p3	p2	p1	c	i3	i2	i1
(MGP-PD)											
27378	0	1+	3								
31493	0	3-	3								
27359 right	1+	3+		3							
31498		3+	3								
31481			3/4								
31497		4-	4								
27377	2	4									
27375	21	4.	4 .	2	1						
27379	37	4+	4+	3+	1						
0 - 0 0 1 1	3+	4+	4+	3+	1						
27382 right	3+	4+	4+	3+	1				1	0	0
27382 right 27382 left	3+	4+	4+	3+	1				1	0	0
27382 right 27382 left 27384 left	3+	4+	4+	3+					1	0 0 2/3	0 0 2/3
27382 right 27382 left 27384 left 27360 right	3+	4+	4+	3+					1	0 0 2/3 3	0 0 2/3 2/3
27382 right 27382 left 27384 left 27360 right 27384 right	3+	4+ 4/5	4+	3+			2/3	2/3	1	0 0 2/3 3	0 0 2/3 2/3 2/3
27382 right 27382 left 27384 left 27360 right 27384 right 31412	3+	4+ 4/5 4	4+ 4/5 4++	3+			2/3	2/3	1	0 0 2/3 3	0 0 2/3 2/3 2/3

Tab. 1 - Wearing stage of *A. monsvialense* teeth. The wear was estimated considering the features of the masticatory surface of each tooth, as explained in "Material and methods".

by a seller (Fig. 3b). Notwithstanding, the species *A. monsvialense* represents today one of the oldest and smallest European *Anthracotherium*, and consequently, it is one of the putative links between the European and Asian species after the Asian radiation.

MATERIAL AND METHODS

Materials: all the described specimens are housed in the Museo di Geologia e Paleontologia, Università di Padova (Padova, Italy), being part of the permanent exhibition of the museum. The collection from Monteviale includes numerous fossils (95 specimens listed below in "Referred Material").

A. monsvialense is here re-described and compared with the specimens of A. magnum, housed at the Università degli Studi di Genova (the whole collection from Cadibona) and in the Museo

di Storia Naturale, Università di Firenze (Section of Geology and Paleontology) (casts, inventory number: IGF 5089V; IGF 11615V; IGF 2107V; IGF 11621V; IGF 4167V). Such casts reproduce real fossils from Cadibona (Savona, northwestern Italy) and are labeled as *Anthracotherium gastaldii* Squinabol, 1890 (synonymized to *A. magnum*; Kotsakis 1986).

For relative wearing stage estimation of the teeth, we numbered the progressive development of visible inner dentine, identified as follow: 0 = tooth not worn but almost partially emerged; 1 =enamel partially worn; 2 = first comparison of dentine in less than half of masticatory surface; 3 = all the cusp/cuspids are worn and show dentine; 4 = dentine of different cusps/cuspids is joined to each other; 5 = the masticatory surface of the tooth is completely worn (Tab. 1). Stage 0 is considered as sub-adult, stages 1-2-3 are considered completely adults and stages 4-5 are senile. The minimal number of individuals (MNI) has been estimated using wear stages of dental series with almost two teeth in sequence.

State of preservation: the fossils from Monteviale show a characteristic state of preservation: post-depositional events and diagenesis concentrated high percentages of pyrite in the bone cavities. Usually, this is dangerous for fossil preservation: high humidity conditions easily start oxidation, leading to the formation of sulphates with a significant increase of the volume of the mineral. Therefore, it causes the chemical modification of the specimen starting from the inner side, with the formation of cracks on the bone surface until the complete and un-reversible destruction of the fossil (Mikuž & Aničič 2007; Del Favero et al. 2012). In order to isolate and protect the fossils from the external environment, during the first half of XX century all the specimens had been painted using a grey varnish, partially covering specific features. Recently, the remains of A. monsvialense have been partially restored to their original status (Maritan 2004-2005), allowing to describe them in great detail. A new step of restoration of the Monteviale's collection has been completed in the spring of 2016.

Measurements and tools: all the measurements were taken with a vernier caliper (resolution of 0.05 mm) using the protocol proposed by Driesch (1976). Statistical test MANOVA has been performed using Past3.02a tools (Hammer & Harper 2001). Pictures have been taken using a Nikon D300S digital reflex camera; colors and contrasts were modified for a better color and lighting effect.

Terminology: taxonomic and morphological nomenclature follows the standard outlined by Boisserie et al. (2010) for bunodont cetartiodactyls, partially modified by Barone (2006) and Lihoreau & Ducrocq (2007). Dental nomenclature of upper and lower molars is shown in Fig. 4.

Anatomical and measurements abbreviations: I/i = upper/lower incisor; C/c = upper/lower canine; P/p = upper/lower premolar; M/m = upper/lower molar; d = upper/lower deciduous tooth, L = length, W = width.

Institutional abbreviations: MGP-PD, Museo di Geologia e Paleontologia, Università degli Studi di Padova; MP-FI Museo di Storia Naturale, Università degli Studi di Firenze; abbreviations reported in the Appendix 1 are from Tsubamoto et al. (2002).

Systematics

Order **Cetartiodactyla** Montgelard, Catzeflys & Douzery, 1997 Family Anthracotheriidae Leidy, 1869 Subfamily Anthracotheriinae Leidy, 1869 Genus *Anthracotherium* Cuvier, 1822



Fig. 4 - Scheme of the dental nomenclature of third upper and lower molars. Other teeth follow the same nomenclature. Ectoparacristule, preparacristule, postprotocrista, endoprotofossa and endometacristule are not shown in figure, but their position can be inferred by the anatomical description of M3 (schemas from Boisserie et al. (2010) applied to anthracotheres).

Anthracotherium monsvialense De Zigno, 1888

Figs 2, 5-19

1865 Anthracotherium magnum Cuv.; von Schauroth: 266, pl. XXX, fig. 2.

1865 Anthracotherium magnum minore (sic) Cuv.; Beggiato: 8, figs 6, 7.

+1888 Anthracotherium monsvialense Zigno: 37, pl. I.

1897 Anthracotherium magnum Cuv.; Flores: 93.

1910a Anthracotherium monsvialense de Zigno; Stehlin: 176, figs 1-2.

1915 Anthracotherium monsvialense Zigno; Fabiani: 308.

1926 Anthracotherium monsvialense De Zigno; Dal Piaz: 55.

1929 Anthracotherium monsvialense; Dal Piaz: 911.

1932 Anthracotherium monsvialense De Zigno; Dal Piaz: 2, figs 1-4 (text figs), pls I-XVI.

1980 Anthracotherium monsvialense; Altichieri: 173 (text fig.).

1986 Anthracotherium monsvialense Zigno; Kotsakis: 143.

1997 Anthracotherium monsvialense De Zigno; Mietto: 134, fig.

2005 Anthracotherium monsvialense Zigno; Kotsakis et al.: 135, fig. 18.

> 2006 Anthracotherium monsvialense De Zigno; Mietto: 41, fig. 25. 2016 Anthracotherium monsvialense Zigno; Pandolfi et al.: 23, figs

Holotype: MGP-PD 26556: palate with upper tooth series (Fig. 3).

Range: Italy, and probably south France, western Switzerland and west-middle Germany, Spain and Turkey.

Referred Material: Splancnocranium: MGP-PD 5740, 26556, 27359, 27360, 27382, 31503; upper canines: MGP-PD 27367, 27369, 31490; upper incisors: MGP-PD 27371, 27372, 27373, 27385, 31477a, 31477c, 31478a, 31489, 27370 (cfr. A. monsvialense); upper premolars and molars: MGP-PD 31410, 31411, 31492, 31496, 31499, 31500, 31506a, 31506b, 31512; upper tooth series: MGP-PD 5548, 27363, 27364, 27366, 27388, 31413, 31486, 31488, 31491, 31495, 31501, 31502; teeth with fragmented mandibles: MGP-PD 27377, 27378, 27379, 27384, 27386, 27387, 31412, 31432, 31479, 31480, 31493, 31507c; lower canines: MGP-PD 27368; lower incisors: MGP-PD 5551, 31415, 31497, 31498; premolars and molars: MGP-PD 5549, 12837, 27376, 27380, 27381, 31414, 31478b, 31482a, 31484, 31494, 31504, 31505, 31513a-c; fragmented teeth: MGP-PD 31477b, 31478c, 31482b, 31483); juvenile deciduous teeth: MGP-PD 27361, 27362, 27383, 31481.

Postcrania are represented by a distal epiphysis of a right humerus (A. monsvialense, MGP-PD 27389), complete, but fragmented hindlimb (MGP-PD 12900); isolated postcranial remains such as a left ulna (MGP-PD 31401), two fragments of tibia (MGP-PD 31507a-b), patella (MGP-PD 27390, 31487), metapods (MGP-PD 27393, 27394, 27395, 27396), tarsals (right astragalus MGP-PD 27391, MGP-PD 31485), and phalanxes (MGP-PD 27397, 27398, 27399).

Description

Upper teeth: molars are well imbricated, low-crowned and bunodont. Among upper teeth there is not a diastema; incisors are more and more horizontally inserted on the premaxillar with the fist one enlarged and completely horizontally oriented (Fig. 5).

M3: the third upper molar shows a sub-rectangular shape with a prevalent protocone. The mesial margin of the *cingulum* has a noticeable mesiostyle, it is not joined to the protocone by the ectoprotocrista. In addition, the mesiostyle ends with a variably developed bulged *cingulum* at the lingual side. The parastyle shows a lingual low crest that reaches the paracone bulge. The lingual side of the tooth is well developed and enlarged, with a double lobed appearance formed by protocone and metaconule (Fig. 6). On all the M3 of the collection but on MGP-PD 27363, entostyles are really low-crowned and un-developed. The dimensions of all M3s, MGP-PD 27363 and MGP-PD 27359 are larger than the other homologous, nevertheless the entostyle is similar in size in all M3. This allows ruling out the hypothesis of a sex-related cause for the development of the entostyle.

A basal notch isolates the distostyle to the postmetacristule, but the distostyle is ideally connected to the metaconule by such crest. Distostyle is

16-18.

130.



Fig. 5 - A. monsvialense MGP-PD 27360: ventral side of upper and lower incisors and canines (A) and left upper incisors series and canine in lateral view (B). Upper incisors are canine-like and obliquely inserted on the jaw. Upper canines are labeled by the white arrow.

not joined to the metacone. Metastyle has the same size of distostyle and it is located on the distolabial side. The postmetacrista connects the metastyle to the metacone on MGP-PD 27363 and MGP-PD 27364 (Fig. 6). On MGP-PD 27366 postmetacrista and metastyle are separated, as well as distostyle and postmetaconule. The ectostyle stands out in the labial margin with a large columnar structure and without any labial crests; it is the only one connected to two tips, the metacone and paracone (respectively by the premetacrista and postparacrista). A real ectocristyle is not visible and the crest begins and diverges directly from the ectostyle. The paracone seems to be quite isolated in comparison to other cusps. It has a conical mesiolabial side, whereas the medial side has a more jagged appearance with a variable number of fossae.

All the upper molars exhibit an accessory cusp, the paraconule, which is a characteristic feature of all anthracotheres. It is located on the trigonid, between the paracone and the protocone, and it is connected to the latter with the preprotocrista. The paraconule has two crests: the ectoparacristule that is distally directed to the inner side of the molar and the preparacristule that links it with the preprotocrista in the upper side. The preparacristule reaches also the mesial cingulum, but it does not reach the labial parastyle. The endoprotocrista is connected to the premetacristule at the centre of the tooth. The postprotocrista is distally directed and completely isolated; it is also separated from the endoprotocrista by a thin endoprotofossa. The premetacristule and the endometacristule origin from the metaconule; the former is high and joined to the endoprotocrista, whereas the latter is transversally directed and shows a bi-lobed shape on MGP-PD 27364. Both are not connected to the metacone and a welldefined sagittal valley divides the two distal cusps. The metacone shows a postmetacrista that develops sub-parallel to the sagittal axis (Fig. 6).

M2: the lingual side is characterized by a less bulged feature of basal protocone and metaconule in correspondence to the third molar. Distostyle and metastyle are not developed on the distal edge of the tooth. Postmetaconule and postmetacone end as simple crists (they appear truncated). The tooth is about 1/3 smaller than M3 (Tab. 2). The ectostyle is developed as well as the homologous of M3, whereas the entostyle is faint.

M1: only MGP-PD 27363 (Fig. 7) preserves useful morphological features. The tooth is morphologically similar to M2. The quadrangular collar is simplified and undeveloped, the crest on the lingua-mesial edge of the tooth is smaller than in other molars, and the ectostyle is pronunced on the lingual side.

P4: the last premolar displays only two cusps on the labial (paracone) and lingual (protocone) si-



Fig. 6 - A. monsvialense MGP-PD 27364: fragment of the palate with right and left M1 (both fragmented), M2 and M3.

<i>p</i> value:	1,82E-21						_
M3	birmanicum	crassum	magnum	monsvialense	pangan	tenuis	
birmanicum		2,31E-03	1,62E-08	5,49E-11	2,07E-09	9,73E-03	
crassum	2,31E-03		9,75E-06	9,54E-07	1,42E-05	3,22E-04	
magnum	1,62E-08	9,75E-06		4,13E-05	5,23E-05	4,10E-06	1
monsvialense	5,49E-11	9,54E-07	4,13E-05		0,3597	1,31E-09	
pangan	2,07E-09	1,42E-05	5,23E-05	0,3597		3,30E-08	
tenuis	9,73E-03	3,22E-04	4,10E-06	1,31E-09	3,30E-08		1
							-
p value:	7,92E-30	7					
m3	birmanicum	crassum	magnum	M-monsvialense	monsvialense	pangan	tenuis
birmanicum		0,05557	2,9014E-10	4,3325E-04	4,8567E-06	1,2108E-02	1,5915E-02
crassum	0,05557		4,6034E-12	2,4579E-04	6,7907E-06	1,8268E-02	3,3715E-07
magnum	2,9014E-10	4,6034E-12		4,6148E-04	7,6273E-06	2,4129E-04	3,4301E-20
M-monsvialense	4,3325E-04	2,4579E-04	4,6148E-04		0,17736	0,53433	1,8187E-09
monsvialense	4,8567E-06	6,7907E-06	7,6273E-06	0,17736		0,19662	1,2349E-13
pangan	1,2108E-02	1,8268E-02	2,4129E-04	0,53433	0,19662		1,0362E-06
tenuis	1,5915E-02	3,3715E-07	3,4301E-20	1,8187E-09	1,2349E-13	1,0362E-06	

Tab. 2 - MANOVA test among upper and lower teeth of the compared anthracotheres. Positive results are reported in bold.

des (Fig. 3). A continuous *cingulum* is slightly developed on the mesial and lingual edges, whereas it is more developed distally, with a low-crowned metastyle and distostyle. Parastyle is on the mesial-labial corner of the tooth. The two cusps are divided by a deep sagittal *fassa* in the middle of the tooth. The protocone is bulged on the lingual side and shows a steep and flat surface on the labial one (formed by the preprotocrista and the distally directed endoprotocrista) (Fig. 7), while the paracone is conical with



Fig. 7 - A. monsvialense: upper P3, P4, and molar series of MGP-PD 27363 on occlusal (A), labial (B) and lingual (C) views. The surface of teeth is coated with ammonium chloride in order to maximize different features. Scale bar 2 cm. See the online content of Appendix 2 for the 3d rendering.

several quite developed crests. Distally, the distoprotocrista is completely worn.

P3: the third premolar is transversally oriented on the upper jaw, and displays a high-crowned paracone. Preparacrista and postectoparacrista reach the basal end of the enamel, on the mesio-lingual and disto-labial edge of the tooth, respectively. On the lingual side, the tooth enlarges distally, for the development of a distal crest with a metastyle (Fig. 3).

P2: P2s are preserved on the holotype (MGP-PD 26556) and on MG-PD 27365. P2 is morphologically similar to P3. The tooth is mono-cuspidate and asymmetrical, for the presence of a mesially directed and high-crowned paracone. The distal side of the tooth ends with a flat surface. There is a postparacrista and a *cingulum* on the distal edge of the tooth.

The distal side of P2 is larger than the mesial one for the presence of the metastyle and a quite developed postparacrista.

P1: the first premolar on MGP-PD 27365 is similar to the P2, but differs by having a smaller size, by lacking a real ectoparacrista (only a small tubercle is present) and by having a less pronounced distal feature.

Upper canines: upper canines are well preserved on MGP-PD 27360 (Fig. 5) and MGP-PD27365 (Fig. 8). They have a conical shape with triangular profile in labial view. MGP-PD27360 has a higher crown in comparison to MGP-PD 27365. This difference is commonly related to the sex of the animals as suggested by Lihoreau & Ducrocq (2007), the former probably representing the male. There is not a real diastema, neither before nor after the canines on MGP-PD 27360, nor on MGP-PD 5548, the other putative females; indeed, a short diastema exist on MGP-PD 27365, between the canine and the last incisor.



Fig. 8 - *A. monsvialense* MGP-PD 27365: fragment of upper dental series with the I3, high-crowned canine, P1 and P2.

Upper incisors: I2 and I3 are canine-like, with a single cusp and a single root, a convex mesial edge and a longer concave distal edge. A developed medial crest runs from the tip to the distal end of the tooth. The enamel-root boundary is oblique and goes down from the mesial side to the distal one. I1s are large and spatular-shaped. Crowns are flat and slightly curved towards the lower jaw, enlarging the enamel surface at the mesial and distal ends of the teeth.

Lower teeth: the cheek teeth are mesio-distally joined to each other and there is not a diastema between molars and premolars, neither among the other teeth of the lower jaw. The trigonid is welldeveloped on m3, and less developed and smaller on m2 and m1. On m1, the trigonid can occupy less than half of the tooth. The premolars, m1, and m2 have two joined roots in correspondence of the trigonid and talonid, whereas m3 has four or five independent roots, under each cuspid and an enlarged one under the distal buldged hypoconulid.

m3: the tooth has a sub-rectangular shape, enlarged on the mesial edge (Figs 9, 10a, b). Mesially to the metaconid, the preentostylid is present. This stylid occurs on all the analyzed m3 and is located above the enamel-root boundary on the mesio-lingual side. On MGP-PD 27378, the postectostylid is little developed and on MGP-PD 27376 (Fig. 10a, b) there are two small postectostylids between the hypoconid and hypoconulid (such teeth are also the largest m3 of the entire collection - Appendix 1).

Premetacristid and postectometacristid are straight and oriented along the sagittal plane of the tooth. Premetacristid is joined to the preprotocristid with multiple (two or three) little tubercles. The entometacristid is represented by a short cristid, which is mesio-medially oriented. Postmetacristid and postprotocristid display a U-shaped wall between the trigonid and talonid. A well-developed prehypocristid divides the transverse fossae in the inner side of the m3 into a labial and a lingual valley. Protoconid and hypoconid are more selenodont than the lingual cuspids; preprotocristid, postprotocristid, prehypocristid and posthypocristid strongly curve toward the lingual side of the tooth and end over the medial plane of the masticatory surface. On the labial side, the above-mentioned cuspids have a bulged neck, forming two developed lobes. The postectoprotocristid is straight oriented and consists of a low cristid on the distal side of the protoconid; it ends in correspondence of the labial transverse valley without a real junction to the prehypocristid, but forming a triangular postprotofossid.

On the talonid, the prehypocristulid is not fused with the posthypocristid, but they join along the inner side of the tooth, at the distal narrowed side of the tooth. The talonid has a bulged neck. The main tip of the talonid, the hypoconulid, has no ectohypocristulid on the labial side. An entoconulid is variably developed on the m3s and usually well developed on the largest specimens. For example, it is a small tip on MGP-PD 27380 (Fig. 10c) and a real



Fig. 9 - A. monsvialense MGP-PD 31434: dorsal (A) and lingual (B) views. Such fossil is the better mandibular horizontal branch preserved with molars (m1-m3).



Fig. 10 - A. monsvialense: lower third molars with different development of entoconulid: large left m3 (MGP-PD 27376) in lingual (A) and dorsal (B) view; smaller right m3 (MGP-PD 27380) on occlusal (C) view.

cuspid on MGP-PD 27376 (Fig. 10a, b), respectively 56.4 mm and 49.8 mm of length. When developed, the entoconulid is well-divided from the other cuspid of the talonid and it has a single crest, a preentocristulid; the latter occurs on the mesial-lingual side reaching the prehypocristulid, but without a real junction. In fact, the medial valley of the talonid is long and continues between the hypoconulid and the entoconulid and the respective cristids.

m2: this tooth is characterized by a smaller size than m3. The absence of a talonid is substituted by a medial distal stilid (distostylid) and by a general simplify feature, with the repetition of cristid morphologies. There is a small preentostilid, a postectoprotocristid and a postectometacristid; the prehypocristid is the most developed cristid of the tooth.

m1: this tooth is labial-lingually thinner and shorter if compared to the m2. It has a simplified morphology. The preentostilid is still present, but there are not other stillds or a clear evidence of a distostylid.

p4: the last lower premolar has a single main protoconid and a distal *cingulum* with the broad-based tubercle of the hypoconid. There is an appreciable metaconid on the lingual side of the protoconid, aligned to the lingual cuspids of the following molars. The wearing erased most of the crests in all the specimens from Monteviale, but the presence of a preprotocristid, a postprotocristid and an



Fig. 11 - A. monsvialense: left mandible (MGP-PD 27375) of A. monsvialense on occlusal (A) and lateral (B) views with p3-m3.

Fig. 12 - *A. monsvialense* MGP-PD 27384: lower jaw with right m1, fragmented p4-p2, p1 (white arrows), canine and spatular arrangement of left incisors.



endometactistid could be inferred (Fig. 11). A distostylid and an entocristid are visible only on the fragmented p4 of MGP-PD 27378. The paracone is not present in the lingual side of the tooth.

p3: the third premolar also has a single protoconid and a distal less developed *cingulum*. The tooth differs from p4 by the presence of two distal tubercles at the lingual and labial side of the tooth; such tubercles emerge at the mid-height of the protoconid crown, the lingual one being larger than the labial tubercle (Fig. 11).

p2: p2 is not preserved.

p1: on MGP-PD 27384 (Fig. 12) the first premolar is located on the inner side of the mandibular ramus, due to compaction and diagenesis. The shape of the tooth is similar to that of the canine, with a more developed concave distal edge, a bulged distal side and a single mesially oriented protoconid. Accessorized cristids are not evident.

Lower canines: canines are similar in shape to the upper homologous, but appear more slender. The single cuspid is mesially oriented and the tooth is obliquely inserted in the jaw. The specimen MGP-PD 27384 (Fig. 12) probably belongs to a female individual, due to the simple shape and its small size (Tab. 2). The two canines on MGP-PD 27382 (Fig. 13) are noticeably larger and higher than the previous one, with a more conical feature. They probably belong to a male individual.

Lower incisors: lower incisors are spatula-like, forming a real spade on the *rostrum*. They are cylindrical and dorso-ventrally compressed from the root to the enamel. i3 (Figs 5 and 12) is shorter and larger than i2 and i1. The latter incisor is the smaller of the series.

Mandibular bone: MGP-PD 31432 is a relatively well-preserved left horizontal *ramus* of *A. monsvialense* with m3-m1 in situ (Fig. 9). The preserved mandible is thin with an enlarged distal side below and behind m3. The height of the horizontal ramus behind m3 is 70 mm. The ventral edge of the mandible curves toward the angular process. Other mandibular fragments do not show useful morphological characters and are not described here.

Postcranial bones: In general, postcranial bones from Monteviale are poorly preserved; fossils suffered the effects of post depositional compaction and diagenesis; they are partially destroyed by pyrite oxidation, damaged and quite fragmented.



Fig. 13 - *A. monsvialense:* dorsal (A) and ventral (B) view of the mandibular rostrum (MGP-PD 27382). Left i1 and i2 have been reconstructed.



Fig. 14 - *A. monsvialense:* schematic drawing of MGP-PD 31401, proximal fragment of a left ulna on frontal (A) and medial (B) views.

The forelimb is represented by a right humeral distal end (MGP-PD 27389). The capitum is medial-laterally short and the trochlea is thin and it ends with an angular edge. The medial epicondyle consists of a flat surface (distal width = 52 mm). The olecranon fossa is deep and quite large (epicondilar distance = 30 mm) and the lateral epicondyle is medio-laterally thinner than the medial one (18 and 26 mm respectively).

The proximal side of a left ulna (MGP-PD 31401, Fig. 14) shows a well-developed incisure on the apex of the large olecranon (medial-lateral width > 20.5 mm; dorso-ventral dept of the olecranon above the anconeal process = 49 mm). The trochlear trough is large and consists of a rounded surface. The diaphysis is broken under the humeral articulation. With the exception of a short junction under the trochlear trough, this ulna was not ossified with the radius.

The hind limb is represented by several bones. MGP-PD 12900 consists of a left leg (Fig. 15),



Fig. 15 - Hindlimb of *A. monsvialense* (MGP-PD 12900) and identification of anatomical elements of the autopodium (cal. = calcaneus; astr. = astragalus; nav. = navicular; cun. = cuneiform; ses. = sesamoid; met. = metatarsus; ph. = phalanxes I, II, and III).

and includes the distal epiphysis of the femur, tibia, tarsals, metatarsals, sesamoids and phalanxes. The bones are fragmented and partially embedded in a black-painted slab: therefore, they can be only partially described and measured. The length of the small cuneiform is 26 mm, whereas the medio-lateral width is 21.5 mm. The length of the calcaneus is 100 mm (from the talon to the distal edge) and the distal width of the astragalus is 37 mm. Medial phalanxes (of the second and third digit) are 42 mm long and the median width is 18.5 mm; lateral phalanxes are about 31.5 mm in length and 8 mm in



Fig. 16 - Partially joined metapods of *A. monsvialense* (MGP-PD 27396 and MGP-PD 27394).

width. In the collection of MGP-PD, other two undetermined lateral phalanxes measure respectively 39.5x10.5 mm (MGP-PD 27397) and 38.5x10 mm (MGP-PD 27398).

The two patella have a diamond feature with rounded angle; they are medio-laterally enlarged (GL = 66.5 and 62.7 mm; GB = 44.2 and 48.6 mm, respectively), with a rather flat surface. The caudal surface is divided by an eccentric bulged wall in two articular surfaces, the medial one being thinner and

with well-recognizable outline. MGP-PD 31487 shows the reorganization of the lateral articular surface, perhaps due to a rearrangement of lateral muscles (due to articular stress or after a collision).

Metapodial bones are represented by two specimens. Such extremities consist of two distal fragments of lateral digits and two partially joined medial metapods of about 10 cm in length (MGP-PD 27393, 27394, 27395, 27396). Unfortunately, the proximal epiphysis are not preserved and their anatomical position (*manus* or *pes*) cannot be inferred (Fig. 16).

MGP-PD 31507a-b consists of a proximal and distal portion of a right tibia. The surface of the proximal epiphysis is large: the medial and lateral articular surfaces are quite externally developed and the lateral intercondylar eminence slightly emerges from the articulation surface (the medial one is not preserved). Cranially, the lateral tuberosity starts from the lateral side of the bone, reaching the large and columnar tibial crest (medial-lateral width = 31mm) in the middle, forming a curved bulge on the medial side (Fig. 17). Such feature leaves a deep tendinean groove on the middle side of the tuberosity, usually visible on extant graviportal herbivores (i.e. hippos and elephants). The passage for the extensor muscle is laterally located and is formed by a thin and transversal valley between the lateral articular surface and the cranial tuberosity.

The distal diaphysis of the tibia MGP-PD 31507 has a compressed triangular transverse section (medial-lateral width > 55 mm, cranial-distal width > 31.6). The distal epiphysis (cranial-distal length = 48 mm) is formed by a single well-developed cranio-distally directed medial coclea; it is parallel to the medial edge of the bone and deep. Distally, the medial malleole is rounded. The dorsal

Fig. 17 - *A. monsvialense*: proximal epiphysis of the right tibia (MGP-PD 31507a); the frontalright view is shown (A) in order to underline the tendinean groove in the middle of the cranial tuberosity. The schematic drawing is on frontal view (B).



being probably found only in a few Rupelian (MP21-23) localities, at Bach, Caylux, Quercy, and Villebramar in France, Weinheim in Germany, and Vaulruz in Switzerland (Kotsakis 1986; Sudre 1995; Becker et al. 2004; Scherler 2011 and reference therein). So far, Monteviale represents the first site where such anthracothere had been described, yielding the richest collections of the species in Europe. As in other cetartiodactyls (Lihoreau & Ducrocq 2007), sexual dimorphism of *A. monsvialense* is recognizable by the canine features, and by the development and morphology of m3. As described

development and morphology of m3. As described in the previous section, massive m3 could belong to males and more slender m3 to females, while the development of entoconulid varies accordingly to the size of the tooth and cannot represent a sexual character (Fig. 10).

Since the early discoveries, anthracotheres have been considered semi-aquatic animals, due to the hippo-like size and body proportions, and their barrel and massive body mass; for A. monsvialense the body mass has been estimated to be probably more than 300 kg by Martinez and Sudre (1995; fig. 6) (calculated using the linear regression for mass estimation applied with reserve; the mathematical function is: $Y = 3.16X^{1.482}$ - where Y is the estimated body mass and X the product between lateral length and medio-lateral width), or about 150-200 kg by Tsubamoto (2014, fig 3A) (both body mass estimations are referred to the astragalus MGP-PD 27391). Anyway, the morphological adaptations to a real amphibian life-style seen for extant hippos were not evolved in A. monsvialense and other Anthracotherium species (Pickford 2008).

The presence of lignite beds in the Monteviale locality attests to a wet habitat. The area was characterized by the presence of humid forests and it was probably located close to a coastal brackish lagoon that was characterized by salinity fluctuations (Pandolfi et al. 2016). The whole fauna from Monteviale was linked to such kind of habitat, because the taphonomic setting, with several skeletal parts still in articulation (i.e. MGP-PD 27360, 27396, 12900 just for anthracotheres), suggests that the carcasses did not undergo transportation and consequently are not allochthonous. To date, new observations and analyses should be performed to support this hypothesis.

Both permanent and milk teeth (Fig. 19; MGP-PD 27383; dM3L =24-30 mm; dM3W=20-

Fig. 18 - A. monsvialense: right astragalus (MGP-PD 27391) on dorsal view.

apophysis is moved directly in correspondence of the distal end of the medial coclea, on the lateral side of the bone. The lateral side is partially broken and formed by a concave and medially developed surface.

The well-preserved astragalus (MGP-PD 27391: GL= 64.5 mm, proximal width = 30.5 mm, distal width = 34.7 mm) shows a proximal trochlea with two parallel articular surfaces and a medial deep valley for the tibial articulation; the lateral side is larger and protrudes more dorso-posteriorly than the medial surface. The distal trochlea is partially deflected and distally protrudes with a zigzag edge, caused by an enlargement of the lateral articular facet. The two sides of the astragalus are not aligned to each other, the proximal one being medially directed (Fig. 18).

Comparison

Anthracotherium monsvialense from Monteviale. Evidences of A. monsvialense are rare in Europe,



Fig. 19 - Juvenile right mandible of *A. monsvialense* (MGP-PD 27383), with the milk canine, premolar and molar. The distal tooth is the m1, almost completely erupted. At the same time, a premolar had been erupting between the deciduous canine (dc) and deciduous first premolar (dp1).



28 mm) are present in the collection of Monteviale (NMI = 18), testifying to the presence of different ontogenetic stages in the deposit (young - MGP-PD 27383, sub-adults, adults and senile, considering the dental exploitation) (Fig. 20). Similarly to other mammals, the first molar is the first tooth to become worn (Tab. 1). In the upper jaw, M2 and M3 became worn later, but such pattern can differ from the canonical exploitation (see Tab. 1, MGP-PD 27366 for example). Premolars follow the trend of M2-M3. Unfortunately, dental series with both incisives and cheek teeth are not preserved; rostral teeth are more spotted-distributed and incisors show a similar wearing stage to each others; as a consequence, wearing pattern cannot be evaluated for these teeth. In the lower jaw, the m1 is the first tooth subjected to wearing, followed by the m2. p4 follows the same trend of m2; otherwise m3 starts to be worn when m1-m2 are in the stage 3. In MGP-PD 27384 (Fig. 12), mesial teeth (from p1 to i1) are 2/3 in stage when the m1 is quite completely worn (stage 4/5). Such dental erosion allows us to infer and reiterate a chewing method focused on cheek teeth, in order to manipulate the food (crush and mash) (Tab. 1).

Anthracotheres from Monteviale were browsing animals, as proven by the low-crowned upper and lower molars; the latter being more buno-selenodont. The position and orientation of incisors allow to compare the diet of the species to recent taxa, such as wild boar, peccary and warthogs. These extant taxa display more bunodont and high-crowned cheek teeth than anthracotheres. Lower incisors are used like a spatula in recent suids in order to pull up rhizomes, roots, tubers and fruits, and a similar habit could be inferred for the anthracotheres.

Fossils of *A. monsvialense* from Monteviale show a quite homogeneous and fixed morphology of permanent teeth. Only the entostyle on M3 (Fig. 7) and entoconulid vary in size. Entoconulid varies according to the large and small m3 (Fig. 10), but it is always recognizable, also on the most worn surface of small third molars.

Morphology and size comparison. The genus Anthracotherium is represented by scarce and scattered remains in Europe. Such situation has led to the proliferation of several species which should be put in synonymy if dental intraspecific variability is taken into account (Lihoreau & Ducrocq 2007). To date, the comparison between *A. monsvialense* and



Fig. 20 - Clustered column chart for the estimation of NMI. Tab. 1 has been critically considered in order to divide subadult (sa), adult (a) and senile (s) dental series. Then, upper and lower teeth have been compared obtaining a NMI of 18.



Fig. 21 - Scatter diagram of M3 (A) and m3 (B), lengths (abscissa) and widths (ordinate) of Asian and European species of the genus Anthracotherium. Linear regression for A. monsuialense and anthracotheres from Monteviale are drawn as continue lines (large and thin trendlines, respectively). See Appendix 1 for measurements.

other close chronologically related species is limited to a few characteristics. The difference between the third upper molars of A. monsvialense from Monteviale considered as a whole and the homologous tooth from Dvérce, as drew by Fejfar (1987; fig. 10.6) and determined as Anthracotherium cf. monsvialense, consists of the presence of a loop in correspondence of the labial edge of the transverse valley in the Czech anthracothere. The anthracothere from Dvérce resembles more "cf. Bothriodon sp." from the upper Eocene Mongolian remains of the Ergilin Dzo Formation (Tsubamoto & Tsogtbaatar 2008), aged at the late Eocene, but it is still similar to the fossils from Monteviale for the size. Moreover, these two European anthracotheres differ from each other by the smaller paracone on M3, both in absolute size and relatively to the paraconule, ruling out that they belong to the same species.

The similitude between *A. monsvialense* and *Anthracotherium alsaticum* Cuvier, 1822 has been underlined by several researchers (i.e. Stehlin 1910a; Sudre 1995), both for their chronology and absolute size of teeth (see below for size comparison). Anyway, differences at the level of incisors (Stehlin 1910a) and the relative less marked difference between length and width of *A. monsvialense* with respect to *A. alsaticum* from Lobsann (France) (Scherler 2011) seem to rule out a possible synonymy (Fig. 21).

Compared to A. illyricum Teller, 1886, from Trbovlje (Slovenia), A. monsvialense (and A. magnum too) has not a bulged style on the distal-labial side of P4 (Hilber 1919, tab. 2-3). Spillmann (1974) established the species *Anthracotherium frehi* from an upper Oligocene site close to Freinberg (North Austria). Such anthracothere is similar in size to *A. monsvialense* and *A. alsaticum* (LM3 = 33.5 mm, WmaxM3 = 43.0 mm), but it shows a constricted talon (metacone and metaconule) on M3 with a noticeable reduction of the development and number of styles (Spillmann 1974, fig. 3); in fact, if compared to *A. monsvialense*, teeth of *A. frehi* appear really simplified.

Compared to *A. magnum*, *A. monvialense* shows a relatively smaller ectostyle and the absence of preectostyle on M3 (and upper molars in general), and it differs from the larger European anthracotheres from Rickenbach (MP29, Switzerland), Mouillach and Raynal (France), and Cadibona (Italy) for the relatively smaller size of the paraconule in respect to the protocone and paracone. Such different characteristics give to *A. monsvialense* a mesiodistally un-compressed upper molar appearance, with similar labial-lingual and mesio-distal lengths. Similarly to *A. magnum*, the postprotocrista does not join the premetacristule.

Ducrocq (1995) recognized a possible phyletic relationship among *Anthracotherium pangan* Pilgrim & Cotter, 1916, *Anthracotherium chaimanei* Ducrocq, 1999 from Krabi and *A. monsvialense* from Monteviale, underlying the affinities within the genus and the progressive evolution of styles on P4-M3, of lingual cristids on lower premolars and the orientation of P3. Indeed, some features seem to exclude a close affinity between *A. magnum* and *A. monsvialense*; the paraconid is not visible on the lingual side of p4 in *A. monsvialense*, but it is developed in *A. magnum* and such lacking feature seems to be better related to the anthracotheres from Pondaung, but not to *A. chaimanei* (Tsubamoto et al. 2002, fig. 6). Moreover, similarly to *A. monsvialense*, anthracotheres from Pondaung show two cuspids variably developed on m3 talonid (Tsubamoto et al. 2002), meanwhile such additional cuspids is always well developed in *A. magnum* from Cadibona.

In general, postcranium of the anthracothere from Monteviale is not well-preserved and cannot be compared with European relatives. The two preserved astragali (MGP-PD 27391 and portion of MGP-PD 12900) show the typical morphology of ancient not-ruminant Cetartiodactyla, characterized by not-aligned proximal and distal trochlea and an enlarged surface for the articulation with the navicular bone (Martinez & Sudre 1995).

Metapods of *A. monsvialense* are thick and heavy built, especially the medial ones, with a reduction of lateral digits, including such species within the Anisodactyla (Kowalevsky 1874) together to the younger *A. bumbachense* Stehlin, 1910a (aged from MP22 to MP25, Scherler et al. 2010). Such distal conformation of the limbs could be an adaptation to its massive weight, since larger medial metapods cope better with larger support surfaces, which is useful also in a swampy environment (Coughlin & Fish 2009).

Teeth size comparison among Asian anthracotheres, *A. magnum*, and *A. monsvialense* from Monteviale and other localities reveals a chaotic distribution of measurements within each species, without finite limits among European species and *A. pangan* (Fig. 21a-b). Spatial overlaps among species can be partially correct considering different dental characteristics and chronology (the Asian species being older than the European anthracothere evidences).

Asian species from the Pondaung Formation have been dated to about 37.2 Ma (Tsubamoto et al. 2002) and have been considered in this analysis for their morphological affinities with European anthracotheres.

Anthracotherium crassum Pilgrim & Cotter, 1916 from Pondaung represents the smallest species in our comparison, having the m3 length within a range of 32.4-39.5 mm (n. 9) (Fig. 21b). A. pangan has



Fig. 22 - Increasing size of teeth among different species of anthracotheres, shown with a radar diagram: a) A. tenuis b) A. crassum c) A. birmanicum d) A. pangan e) A. alsaticum f) A. monsvialense from Monteviale g) A. monsvialense from other European sites h) A. magnum i) A. frehi.

a diversified distribution through all the groups, and it is noticeable that it matches with the size of *A. monsvialense*. *A. monsvialense* has an m3 with a length varying within a range of 43.1-56.4 mm (n. 10), whereas the largest European anthracothere, *A. magnum*, occupies the right side of the figure for its massive features and larger teeth (range of 52.5-83.5 mm of length, n. 26). *A. alsaticum* from Ruffels de Giroussens (France; Astre 1927) falls within the largest *A. crassum* variability.

The length-width scatter diagram of M3 (Fig. 21a) shows a partially different situation. All the species are within a rigorous trend of increasing size and *A. pangan* and *A. monsvialense* are completely overlapping in size.

In order to verify such correlations, we applied the statistic test MANOVA to both upper and lower third molars (Tab. 2). Assuming that teeth length and width are normally distributed, the null hypothesis have been tested, consisting on the preconception that all samples are taken from populations with equal multivariate means (*A. alsaticum*, *A. illyricum* and *A. frehi* had been not considered because they are represented by only one tooth each) (Hammer & Harper 2010). The *p* values result lower than 0.05 for both upper and lower third molars. The pairwise comparison reveals no correlation among the species, with the exception of low correlation between *A. crassum-A. birmanicum* for the m3. Indeed, strong correlation results for *A. monsvialense-A. pangan*, underlying their putative affinity.

Combined distribution of measurements in the multi-dimensional chart (Fig. 22) shows again the intrinsic equivalence of size of *A. pangan* (d) and *A. monsvialense* from Monteviale (f). Asian anthracotheres have short third molars, whereas the average measurements of *A. magnum* shift outward (h). The trend supports an exponential increase of size when the underlying bone (maxillar and mandible) allows the elongation and widening of the tooth.

DISCUSSION

The faunal assemblage and chronology of Monteviale has been recently analysed and discussed by Pandolfi et al. (2016), referring such fauna to the lowermost Rupelian, MP21. The mammal fauna recorded at Monteviale is, as a whole, related to the migration event that follows the *Grande Coupure*. It settled in the area of Monteviale inhabiting a fresh water humid environment, given the presence of frogs (genus *Palaeobatrachus*). At the same time, the presence of Dugongidae in the site testifies to marine transgression into a coastal brackish lagoon environment (Pandolfi et al. 2016).

Among the 15 taxa described at Monteviale, *A. monsvialense* is one of the largest species along with *Epiceratherium*, and it represents the most significant species for dating the site, because its presence in Europe is restricted to the first part of Oligocene (MP21-MP23).

While the Monteviale's fauna includes several endemic species (Pandolfi et al. 2016), *A. monsvialense* can be somehow considered a tile of a different puzzle, being recorded in other European sites. In fact, the European invasion from Asia corresponds to the acquisition of a larger size for anthracotheres.

Our results support Ducrocq's (1999) conclusions about an evolutive pattern of anthracotheres, with a progressive radiation of new species; from the smaller Asian specimen, we attend to the evolution of midium-sized lower Oligocene animals (including *A. pangan* and *A. monsvialense*) and finally to the largest *A. magnum* of MP25-MP29 in Europe.

Along with A. monsvialense, the collection of Anthracotheridae from Monteviale includes a single m3 (MGP-PD 14757) reported as "Anthracotherium (Prominatherium) dalmatinum" in the inventory label, and a small amount of fossils attributed to the littleknown Anthracochoerus fabiani Dal Piaz, 1931, and Anthracochoerus stehlini Dal Piaz, 1931. Such fossils have been recently restored and an ongoing study will give us the opportunity of a new research about their morphological features (Ghezzo et al. in preparation). At the current state of the art, we can only underline that the presence of A. (Prominatherium) dalmatinum, if confirmed, could represent the second Italian occurrence of such species along with the teeth from Grancona (Bona & Grandi 2014).

CONCLUSIONS

Geological evidence suggests an insular environment for the area around Monteviale and the northestern Venetian region during the earliest Oligocene, also supported by the fauna discovered in the lignite deposits of Monteviale and by the presence of several endemic species (Kotsakis 1986; Pandolfi et al. 2016).

Our analysis reveals a wider context scenario, retracing the anthracotheres discoveries in the Italian Peninsula. The putative ancestors of *A. monsvialense*, intermediate in size between the late Eocene and late Oligocene species, probably reached Europe across an island-chain formed by the current Iran, Anatoly and the Balkans (Ducrocq 1995; Böhme et al. 2014, Pandolfi et al. 2016), populating the area of Monteviale. The older Anthracotheridae finding at Dvérce (Czech Republic) testifies to a possible migration wave of the Asian *Anthracotherium* also towards North and East Europe.

A. monsvialense is a distinct and valid species recognized since the second half of XIX century. The morphological differences with the younger A. magnum confirm that the two species shared the same phyletic branch, the lifestyle and a subtropical habitat, but at different times. Probably, the geographical range of A. monsvialense was more related to islands and relative southern latitudes with respect to that A. magnum, and its chronological distribution in Europe was shorter, starting at about the "Grande Coupure" event and completely disappearing around 30 Myr (late early Oligocene).

Our results strongly suggest a progressive larger size acquired by anthracotheres in about 5 Ma, from the late Eocene to the late Oligocene and during the Asia-European migration.

Furthermore, the new description here provided opens new possibilities for morphological comparison among Eurasian anthracotheres, and it emphasizes the importance of anthracotheres in the Alpine High context.

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References

- Abel O. (1910) Kritische Untersuchungen über die paläogenen Rhinocerotiden Europas. Abh. k. k. geol. Reichsanst. Wien, 20(3): 1-52.
- Accordi B. (1951) Resti di antracoterio nell'Oligocene di Chiuppano (Vicenza). Ann. Univ. Ferrara, n. Ser. Sez. IX - Sci. geol. paleont., 1(1): 1-36.
- Agassiz L. (1832) Untersuchungen über die fossilen Süßwasser-Fische der tertiäre Formationen. *Jb. Mineral., Geogn., Geol. Petrefaktenkd.*, 3: 129-138.
- Agassiz L. (1842) Nomenclator Zoologicus, continens nomina systematica generum Animalium tan viventium quam fossilium, secundum ordinem alphabeticum disposita, adjectis auctoribus, libris, in quibus reperiuntur, anno editionis, etymologia et familis, ad quas pertinent, in singulis classibus. Fasciculus I continens Mammalia, Echinodermata et Acalephas. Sumptibus et typis Jent et Gassman, Soloduri. Various pagination.
- Agnasson I. & May-Collado L. (2008) The phylogeny of Cetartiodactyla: the importance of dense taxon sampling, missing data, and the remarkable promise of cytochrome b to provide reliable species-level phylogenies. *Mol. Phylogenet. Evol.*, 48: 964-985.

- Aguilar J.-P., Legendre S. & Michaux J. (1997) Actes du Congrès BiochroM'97. *Mém. Trav. E.P.H.E. Inst. Montpellier*, 21: 769-805.
- Altichieri L. (1980) Il giacimento di Monteviale (Vicenza). In: AAVV - I vertebrati fossili italiani. Catalogo della Mostra Verona, 1980: 173-174. Tipografica "La Grafica", Vago (Verona).
- Antoine P-O., Métais G., Orliac M.J., Peigné S., Rafaÿ S., Solé F. & Vianey-Liaud M. (2011) - A new late Early Oligocene vertebrate fauna from Moissac, South-West France. C. R. Palevol, 10: 239-350.
- Aubry M.-P., Ouda K., Dupuis C., Berggren W.A., Van Couvering J.A., Ali J., Brinkhuis H., Gingerich P.R., Heilmann-Clausen C., Hooker J., Kent D.V., King C., Knox R.W.O.B., Laga P., Molina E., Schmitz B., Steurbaut E. & Ward D.R. (2007) The Global Standard Stratotype-Section and Point (GSSP) for the base of the Eocene Series in the Dababiya section (Egypt). *Episodes*, 30: 271-286.
- Barone R. (2006) Anatomia comparata dei mammiferi domestici. Volume 1: Osteologia. Il Sole 24 Ore Edagricole.
- Beggiato F.S. (1865) Anthracotherio di Zovencedo e Monteviale. Mem. Soc. ital. Sci. nat. Milano 1: 1-10, 1 pl.
- Becker D., Lapaire F., Picot L., Engesser B. & Berger J-P. (2004) - Biostratigraphie et paléoécologie du gisement à vertébrés de La Beuchille (Oligocène, Jura, Suisse). *Rev. Paléobiol.*, vol. spéc. 9: 179.
- Bergounioux F.M. (1954) Les cheloniens fossiles des terrains tertiaires de la Venetie. Mem. Ist. Geol. Min. Univ. Padova, 18: 1-115.
- Blondel C. (2001) The Eocene-Oligocene ungulates from Western Europe and their environment. *Palaeogeogr., Pal-aeoclimatol.*, *Palaeoecol.*, 168: 125-139.
- Boisserie J-R., Lihoreau F., Orliac M., Fischer R.E., Weston E.M. & Ducrocq S. (2010) - Morphology and phylogenetic relationships of the earliest known hippopotamids (Cetartiodactyla, Hippopotanidae, Kenyapotaminae). Zool. J. Linnean Soc., 158: 325-366.
- Boisserie J.R., Fischer R., Lihoreau F. & Weston E. (2011) -Evolving between land and water: key questions on the emergene and history of the Hippopotamidae (Hippopotamoidea, Cetancodonta, Cetartiondatyla). *Biol. Rev. Cambridge Phil. Soc.*, 86: 601-625.
- Bona F. & Grandi F. (2014) The geologically oldest terrestrial mammal of the Italian peninsula. XII EAVP, Abstract book: 30.
- Borson É. (1820) Note sur des dents du grand mastodonte trouvées en Piédmont et sur des machoides et dens fossiles prises dans la Mine de Houille de Cadibona Proghe Savone. Mem. r. Accad. Sci. Torino (Mem. Cl. Sci. fis. nat.), 27: 31-42.
- Bowen G.J., Koch P.L., Gingerich P.D., Norris R.D., Bains S. & Corfield R.M. (2001) - Refined isotope stratigraphy across the continental Paleocene-Eocene boundary on Polecat Bench in the northern Bighorn Basin. In: Gingerich P.D. (Ed.) - Paleocene-Eocene stratigraphy and biotic change in the Bighorn and Clarks Fork Bains, Wyoming. Univ. Michigan Pap. Paleont., 33: 73-88.

- Bruijn H. de & Rümke C.G. (1974) On a peculiar mammalian association from the Miocene of Oschiri (Sardinia). I-II. Proc. k. nederl. Akad. Wetensch., B, 77: 44-79.
- Coughlin B.L. & Fish S.E. (2009) Hippopotamus Underwater Locomotion: Reduced-Gravity Movements for a Massive Mammal. J. Mammal., 90: 675-679
- Cuvier G. (1822) Recherches sur les ossements fossiles, où l'on rétablit les caractères de plusieurs animaux, dont les révolutions du globe ont détruit les espèces. Tome 5, E. d'Ocagne, Paris.
- Dal Piaz GB. (1926) Osservazioni sulla formula dentaria del genere *Anthracotherium*. *Atti Acc. Sci. Veneto-Trentino-Istriana*, Ser. 3, 17: 54-60.
- Dal Piaz GB. (1929) Nuove osservazioni sull'Oligocene italiano. Rend. Accad. Lincei, 6(9): 910-913.
- Dal Piaz GB (1930) I mammiferi dell'Oligocene veneto. Propalaeochoerus paronae n. sp. Mem. Ist. Geol. r. Univ. Padova, 8: 1-14.
- Dal Piaz GB. (1931) I mammiferi dell'Oligocene veneto: Anthracochoerus n.g., A. stehlini n. sp., A. fabianii n.sp. Mem. Ist. Geol. r. Univ. Padova, 9: 1-36.
- Dal Piaz GB. (1932) I mammiferi dell'Oligocene Veneto: Anthracotherium monsvialense. Mem. Ist. Geol. r. Univ. Padova, 10: 1-63.
- De Zigno A. (1888) Antracoterio di Monteviale. Mem. Ist. Veneto Sci., Lett. Arti Venezia, 23: 1-35.
- Del Favero L., Fornasiero M.G., Reggiani P., Zorzi F. & Molin G. (2012) - Il restauro dei vegetali fossili esposti nella "Sala delle Palme" del Museo di Geologia e Paleontologia dell'Università di Padova. *Museol. sci., nuova serie*, 6(1-2): 49-57.
- Driesch A. von den (1976) A guide to the measurement of animal bones from archaeological sites. Peabody Museum Bulletin 1. Peabody Museum of Archeology and Ethnology. Harvard University.
- Ducrocq S. (1995) The contribution of Paleogene anthracotheriid artiodactyls in the paleobiogeograpical history of southern Europe. N. Jb. Geol. Paläont., Mh., 6: 355-362.
- Ducrocq S. (1999) The Late Eocene Anthracotheriidae (Mammalia, Artiodactyla) from Thailand. Palaeontographica Abt. A, 252(4-6): 93-140.
- Ducrocq S., Coiffait B., Coiffait P-E., Mahboubi M. & Jaeger J-J. (2001) - The Miocene Anthracotheriidae (Artiodactyla, Mammalia) from the Nementcha, eastern Algeria. *N. Jb. Geol. Paläont., Mb.*, 3: 145-156.
- Ducrocq S., Benammi M., Chavasseau O., Chaimanee Y., Suraprasit K., dong Pha P., le Phuong V., van Phach P. & Jaeger J. (2015) - New anthracotheres (Cetartiodactyla, Mammalia) from the Paleogene of northeastern Vietnam: biochronological implications. J. Vert. Paleontol. (JVP), 35(3), e929139, doi: 10.1080/02724634.2014.929139
- Esu D. & Kotsakis T. (1983) Les vertébrés et les mollusques continentaux du Tertiaire de la Sardaigne: paléobiogéographie et biostratigraphie. *Geol. Romana*, 22: 177-204.
- Fejfar O. (1987) A Lower Oligocene mammalian fauna from Dětaň and Dvérce NW Bohemia, Czechslocakia. Münchner Geowiss. Abh., A, 10: 253-264.

Fejfar O. & Kaiser T.M. (2005) - Insect bone-modification and

paleoecology of Oligocene mammal-bearing sites in the Doupov mountains, northwestern Bohemia. *Palaeontol. Electronica*, 8: 1-11.

- Flores E. (1897) Sul sistema dentario del genere Antracotherium Cuv. Boll. Soc. geol. ital., 16(1): 92-96.
- Főzy I. & Szente I. (2014) Fossils of the Carpathian Region. Gareth Dyke (English text Ed.): 282-284.
- Frost S.H. (1981) Oligocene reef coral biofacies of the Vicentin, northern Italy. In: Toomey D.F. (Ed.) - European fossil reef models. *SEPM spec. Publ.*, 30: 483–539.
- Gingerich P.D. (2006) Environment and evolution through the Paleocene-Eocene thermal maximum. *Trends Ecol. Evol., Paleontol. Ser.*, 21(5): 246-253.
- Gray J. E. (1821) On the natural arrangement of vertebrose animals. *LMR*, 15: 296-310.
- Hammer Ø. & Harper D. (2001) PAST: Paleontological Statistics software package for education and data analysis. *Palaeontol. Electronica*, 4(1), 9 pp.
- Hammer Ø. & Harper D. (2010) Paleontological data analysis. Blackwell Publishing, Blackwell, 351 pp.
- Hellmund M. (1991) Revision der europänischen Species der Gattung *Elomeryx* Marsch, 1894 (Anthracotheriidae, Artiodactyla, Mammalia)-odontologische Untersuchingen. *Palaeontographica Abt.* A, 220(1-3): 1-101.
- Hilber V. (1919) Anthracotherienzähne aus Trifail. *Mitt. natw. Ver. Steiermark*, 55: 107-120.
- Holroyd P.A., Lihoreau F., Gunnell G.F. & Miller E.R. (2010) -Anthracotheriidae. In: Werdelin L. & Sanders W.J. (Eds)
 - Cenozoic Mammals of Africa: 843-851. University of California Press, Berkeley.
- Hooker J.J. (2015) A two-phase Mammalian Dispersal Event across the Paleocene-Eocene transition. *Newsl. Stratigr.*, 48(2): 201-220.
- Janis C.M. (1995) Correlation between craniodental morphology and feeding behavior in ungulates: reciprocal illumination between living and fossil taxa. In: Thomason J.J. (Ed.) - Functional morphology in vertebrate paleontology: 76-98. Cambridge University Press.
- Kelley S. (2002) Excess argon in K-Ar and Ar-Ar geochronology. Chem. Geol., 188: 1-22.
- Koch P.L., Zachos J.C. & Gingerich P.D. (1992) Correlation between isotope records in marine and continental carbon reservoirs near the Palaeocene-Eocene boundary. *Nature*, 358: 319-322.
- Kotsakis T. (1986) Problemi paleogeografici dei Mammiferi fossili italiani: le faune oligoceniche. *Geol. Romana*, 23 (1984): 141-155.
- Kotsakis T., Argenti P., Barisone G., Delfino M., Palombo M.R., Pavia M. & Piras P. (2005) - Il Paleogene. I vertebrati continentali. In: Bonfiglio L. (Ed.) - Paleontologia dei Vertebrati in Italia. Evoluzione biologica, significato ambientale e paleogeografia. *Mem. Mus. cin. di Verona.* 2 *serie. Sez. Sci. Terra*, 6: 131-139.
- Kostopoulos D.S., Koufos G.D. & Christanis K. (2012) On some anthracotheriid (Artiodactyla, Mammalia) remains from northern Greece: comments on the palaeozoogeography and phylogeny of *Elomeryx. Swiss J. Paleontol.*, 131: 303-315.

- Kowalevsky W. (1873) Monographie der Gattung *Anthracotherium* Cuvier und Versuch einer natürlichen Classification der fossilen Hufthiere. *Palaeontographica*, 3(22): 347 pp.
- Kron D. G. & Manning E. (1998) Anthracotheriidae. In: Janis C. M., Scott K. M., and Jacobs L. L. (Eds) - Evolution of Tertiary Mammals of North America. Volume 1: Terrestrial Carnivores, Ungulates, and Ungulatelike Mammals: 381-388. Cambridge University Press, Cambridge, U.K.
- Legendre S. (1995) Le Garouillas et les sites contemporains (Oligocene, MP 25) des Phosphorites du Quercy (Lot, Tarn-et-Garonne, France) et leurs faunes de vertébrès. 14. Paléoécologie. *Palaeontographica Abt. A*, 236: 327-343.
- Leidy J. (1869) The extinct mammalian fauna of Dakota and Nebraska, including an account of some allied forms from other localities, together with a synopsis of the mammalian remains of North America. J. Acad. nat. Sci. Philadelphia, 7 (Ser. 2): 1-472.
- Leonardi P. (1950) Segnalazione di resti di Antracoterio nelle ligniti dei dintorni di Chiuppano (Vicenza). *Rend. Accad. Lincei. Cl. Sci. fis. mat. mat., Ser. VIII*, 9: 360-361.
- Lihoreau F. & Ducrocq S. (2007) Family Anthracotheriidae. In: Prothero D. R. & Foss S. E. (Eds) - The evolution of artiodactyls: 89-105. The John Hopkins University Press, Baltimore.
- Lihoreau F., Blondel C., Barry J. & Brunet M. (2004) A new species of the genus *Microbunodon* (Anthracotheriidae, Artiodactyla) from the Miocene of Pakistan: genus revision, phylogenetic relationships and palaeobiogeography. *Zool. Scripta*, 33: 97-115.
- Lihoreau F., Boisserie J-R., Manthi F.K. & Ducrocq S. (2015) - Hippos stem from the longest sequence of terrestrial cetartiodactyl evolution in Africa. *Nature Communic.*, 6: 6264 doi:10.1038/ncomms7264.
- Made J. van der (1999) Biogeography and stratigraphy of the Mio-Pleistocene mammals of Sardinia and the description of some fossils. In: Reumer J.W.F. & De Vos J (Eds)
 - Elephants have a snorkel! Papers in honour of Paul Y. Sondaar. *Deinsea*, 7: 337-360.
- Made J. van der (2008) New endemic large mammals from the Lower Miocene of Oschiri (Sardinia): observations on evolution in insular environment. *Quatern. Int.*, 182: 116-134.
- Maritan M. (2004-2005) I mammiferi delle ligniti dell'Oligocene di Monteviale (Vicenza) conservati nel museo di Geologia e Paleontologia dell'Università di Padova. Elaborato di Laurea di primo livello in Scienze e Tecnologie per la Natura, Università degli Studi di Padova, 53 pp.
- Martinez J-N. & Sudre J. (1995) The astragalus of Paleogene artiodactyls: comparative morphology, variability and prediction of body mass. *Lethaia*, 28: 197-209.
- Meinolf H. & Kurt H. (1994) Neuere Funde von Prominatherium dalmatinum (H.v. Meyer, 1854) (Artiodactyla, Mammalia) aus dem Eozan von Dalmatien. Mitt. bayer. Staatssamml. Paläont. hist. Geol. 34: 273-281.
- Mennecart B., Scherler L., Hiard F., Becker D. & Berger J-P. (2012) - Large mammals from Rickenbach (Switzerland,

reference locality MP29, Late Oligocene): biostratigraphic and palaeoenvironmental implications. *Swiss J. Palaeontol.*, 131: 161-181.

- Meschinelli L. (1903) Un nuovo Chirottero fossile (Archeopterus transiens, Mesch.) delle ligniti di Monteviale. Atti r. Ist. veneto Sci., Lett. Arti, 62(2): 1329-1344.
- Mietto P. (1988) Aspetti geologici dei Monti Berici. In: AAVV - I Colli Berici, natura e civiltà: 13-23. Signum Edizioni, Padova.
- Mietto P. (1997) I giacimenti lignitiferi oligocenici (Monteviale e Zovencedo). In: Preto D. & Tescari G. (Eds) Solo a Vicenza. Gli endemismi della Provincia. I fossili, le piante, gli animali presenti solo nel territorio vicentino: 134-140. Blended Editrice, Vicenza.
- Mietto P. (2006) La Geologia di Monteviale e le miniere di lignite. Comune di Monteviale (Ed.), 125 pp.
- Mikuž V. & Aničič B. (2007) Zob antrakoterija iz oligocenskih (egerijskih) skladov premogovnika Senovo. *Geologija*, 50(2): 285-292.
- Miller E.R., Gunnell G.F., Gawas M.A., Hamdan M., El-Barkooky A.N., Clements M.T. & Hassan S.M. (2007)
 Anthracotheres from Wadi Moghra, early Miocene, Egypt. J. Paleontol., 88(5): 967-981.
- Montagna C. (1857) Giacitura e condizioni del terreno carbonifero di Agnana e dintorni: 1-167, Napoli.
- Montgelard C., Catzejfis F. M. & Douzeryt E. (1997) Phylogenetic relationships of Artiodactyls and Cetaceans as deduced from the comparison of Cytochrome *b* and 12S rRNA mitochondrial sequences. *Mol. Biol. Evol.*, 14(5): 550-559.
- Ozansoy F. (1962) Les anthracothériens de l'Oligocène inférieur de la Thrace orientale (Turquie). *Bull. miner. Res. Explor. Inst. Turkey*, 58: 85-96.
- Pandolfi L., Carnevale G., Costeur L., Del Favero L., Fornasiero Mg., Ghezzo E., Maiorino L., Mietto P., Piras P., Rook L., Sansalone G. & Kotsakis T. (2016) - Reassessing the earliest Oligocene vertebrate assemblage of Monteviale (Vicenza, Italy). *J. Syst. Palaeont*. http://dx.doi.org/10.10 80/14772019.2016.1147170.
- Pickford M. (2008) The myth of the hippo-like anthracothere: The eternal problem of homology and convergence. *Rev. españ. Paleontol.*, 23: 31-90.
- Pilgrim G.E. & Cotter G.P. (1916) Some newly discovered Eocene mammals from Burma. Rec. geol. Surv. India, 47: 42-77.
- Prothero D.R. & Foss S.E. (2007) The Evolution of Artiodactyls. Johns Hopkins University Press, Baltimore, 384 pp.
- Pomel A. (1847) Note sur les animaux fossiles découverts dans le département de l'Allier. *Bull. Soc. géol. France*, 2(4): 378-385.
- Rasser MW., Harzhauser M. (coordinators) et al. (2008) Paleogene and Neogene of Central Europe. In: McCann T. (Ed.) - The Geology of Central Europe. Volume 2: Mesozoic and Cenozoic: 1031-1140. Geological Society, London.
- Rincon A.F., Bloch J.I., Macfadden B.J. & Jaramillo C.A. (2013) - First Central American record of Anthracotheriidae

(Mammalia, Bothriodontinae) from the early Miocene of Panama. J. Vert. Paleontol. (JVP), 33(2): 421-433.

- Roghi G., Fornasiero M. & Lonigo A. (2013) Achille De Zigno. Impegno civico e attività scientifica. Atti del Convegno per il bicentenario della nascita di Achille De Zigno: 1813-2013, 11 Aprile 2013. Padova University Press, 178 pp.
- Rose K.D. (1996) On the origin of the order Artiodactyla. Proc. natl. Acad. Sci. U.S.A (PNAS), 93: 1705-1709.
- Schauroth K. F. von (1865) Verzeichnis der Versteinerungen im Herzoglich. Naturaliencabinet zu Coburg, XVI, 328 pp.
- Scherler L. (2011) Terrestrial palaeoecosystems of large mammals (Tapiridae, Anthracotheriidae, Suoidea) from the Early Oligocene to the Early Miocene in the Swiss Molasse Basin: biostratigraphy, biogeochemistry, palaeobiogeography and palaeoecology. PhD Thesis, University of Fribourg, 213 pp.
- Scherler L., Becker D. & Berger J-P. (2010) A Tooth tale: the Anthracotherium bumbachense (Artiodactyla, Mammalia) from Beuchille (Delémont, Switzerland). 8th Swiss Geoscience Meeting, Fribourg.
- Scherler L., Mennecart B., Hiard F. & Becker D. (2013) Evolutionary history of hoofed mammals during the Oligocene-Miocene transition in Western Europe. *Swiss J. Geosci.*, 106(2): 349-369.
- Sieber R. (1935) Osservazioni preliminari sui resti di Anthracotherium magnum Cuvier di Cadibona conservati nell'Istituto Geologico di Torino e nel Museo Civico di Savona. Rin. It. Paleontol., 41: 72-74.
- Sileem A.H. & Hewaidy A.G.A. (2015) New postcranial specimens of the Anthracotheriidae (Mammalia; Artiodactyla) from the Paleogene of Fayum Depression, Egypt. *Int. J. sc. Eng. appl. Sci. (IJSEAS*), 1(8). ISSN: 2395-3470.
- Smith R., Smith T. & Sudre J. (1996) Diacodexis gigasei n.sp., le plus ancien Artiodactyle (Mammalia) belge, proche de la limite Paléocène-Eocène. Bull. Inst. r. Sci. nat. Belg., Sci. Terre, 66: 177-186.

- Soe A.N. (2008) A new study of the anthracotheres (Mammalia, Artiodactyla) from Pondaung Formation, Myanmar: systematics implications. *Palaeovertebrata*, 36(1-4): 89-157.
- Spillmann F. (1974) Ein neues Anthracotheirum aus den Oberoligozänen linzer sanden (Anthracotherium frehi n.sp.). Jh. Oberösterreichischen Musealvereines: 179-198.
- Squinabol S. (1890a) Cenni preliminari sopra un cranio ed altre ossa di Anthracotherium magnum Cuv. di Cadibona. Estratto dagli Atti Soc. lig. Sci. nat. geogr., 1(1), 11 pp.
- Squinabol S. (1890b) Rivista dei grossi Anthracotherium di Cadibona. Boll. Soc. geol. ital., 9: 515-571.
- Stehlin H.G. (1910a) Zur revision der europäischen Anthracotherien. Verh. natf. Ges., Basel, 21: 165-185.
- Stehlin H. G. (1910b) Remarques sur les faunules de Mammifères des couches éocènes et oligocènes du Bassin de Paris. Bull. Soc. géol. France, 9(4): 488-520.
- Sudre J. (1995) Le Garouillas et les sites contemporains (Oligocene, MP 25) des Phosphorites du Quercy (Lot, Tarnet-Garonne, France) et leurs faunes de vertebres. 12. Artiodactyles. *Palaeontogr. Abt.* A, 236: 205-256.
- Tsubamoto T. (2014) Estimating body mass from the astragalus in mammals. *Acta Palaeont. Pol.*, 59(2): 259-265.
- Tsubamoto T., Takai M., Egi N., Shigehara N., Tun S. T., Aung A. K., Soe A. N. & Thein T. (2002) - The Anthracotheriidae (Mammalia; Artiodactyla) from the Eocene Pondaung Formation (Myanmar) and comments on some other anthracotheres from the Eocene of Asia. *Paleont. Res.*, 6(4): 363-384.
- Tsubamoto T. & Tsogtbaatar K. (2008) New specimens of anthracotheriid artiodactyls from the upper Eocene Ergilin Dzo Formation of Mongolia. *Paleontol. Res.*, 12(4): 371-386.
- Ungaro S. (1978) L'Oligocene dei Colli Berici. Riv. It. Paleontol. Strat., 84(1): 199-278.
- Van Valen L. (1966) Deltatheridia: a new order of Mammals. Bull. amer. Mus. nat. Hist., 132: 1-126.