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GEOMORPHOLOGY

The geomorphic history of the Ainoura plain, Kyushu, Japan, based on excavation of the Monzen ruins

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ABSTRACT

The complex coastline of north-western Kyushu, Japan, consisting of cliffs, steep-sided inlets and archipelagos, has resulted from subsidence in mountainous terrain; the region's populated river valleys, having narrow alluvial plains, often suffer rock avalanches. The Ainoura plain is one such alluvial plain. Excavation of the Monzen ruins on the Ainoura Plain by the Nagasaki Prefectural Board of Education has revealed the plain's geomorphic history regarding how the gravels and associated muds and sands were deposited during the last several thousand years. Regarding alluvial succession, mud and sand first filled in the trough, possibly as part of the transgression caused by the Holocene climatic optimum, and gravel later began to replace parts of the sand and mud layers by successive erosion and sedimentation, thereby raising riverbed level. It was confirmed that such conglomerate from the Ainoura plain was derived from the Hokusho landslide area in the hinterland because there is no conglomerate layer in the Hino plain (a small plain southeast of the Ainoura plain). The Ainoura River's thalweg shows that the Monzen ruins are located at a sudden reduction in its gradient, thereby accounting for the thickness of coarse clastic material in that area. The reconstructed Ainoura plain's paleoshoreline during the Holocene climatic optimum is also reported. The paleoshoreline did not move much until people began to reclaim land; it would thus be expected that the similar deposition would happened on the Ainoura plain inside this paleoshoreline.

RESUMEN

La compleja línea de costa en la región noroccidental de Kyushu, Japón, está compuesta por acantilados, ensenadas de laderas inclinadas y archipiélagos, y es producto de la subsidencia de un cinturón montañoso. Los valles aluviales y las estrechas llanuras de inundación de esta región se encuentran densamente pobladas y están sujetas a continuos movimientos en masa. Ainoura es un ejemplo de esta llanura de inundación. La excavación de Monzen Ruin realizada por "Nagasaki Prefectural Board of Education" en la llanura de inundación Ainoura revela la historia geomorfológica; la depositacion de gravas, arenas y lodos en los últimos milenios. En la sucesión aluvial, los primeros en llenar el canal fueron los depósitos de lodos y arenas, probablemente como resultado de la transgresión del óptimo climático que ocurrió durante el Holoceno. Consecutivamente, los depósitos de grava empezaron a reemplazar las capas de arena y lodo mediante la erosión y depositación sucesiva, elevando el lecho del rio. Nosotros corroboramos que el conglomerado de la llanura de inundación Ainoura proviene del deslizamiento de tierras de la zona interior de Hokusho, ya que la Llanura de inundación Hino, ubicada al sureste de Ainoura, no contiene capas de conglomerados. El lecho del rio Ainoura muestra que el Monzen Ruin se localiza en la abrupta reducción del gradiente, lo cual explica el material clástico grueso en esta área. Adicionalmente, nosotros reportamos la paleo-línea de costa de la llanura de inundación de Ainoura durante el optimo climático que ocurrió durante el Holoceno. La paleolínea de costa no se movió significativamente hasta que la población empezó a recuperar tierra. Por consiguiente, nosotros predecimos que el mismo proceso de depositacion podría ocurrir en la llanura de inundación de Ainoura al interior de la línea de costa.

Keywords: alluvial plain, Hokusyo landslide, Kitamatsuura basalt, Kyushu, transgression.

Palabras claves: Basalto de Kitamatsuura, Corrimiento de tierra de Hokusyo, llanura de inundación, Kyushu, Transgreso

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Introduction

The Japanese archipelago is geologically active and therefore suffers from geological disasters; millions of people still live on the islands, coping with disasters. Rock avalanches forming part of the process of sedimentation represent an example of such disasters. The Hokusho landslide area is located in north-western Kyushu (Noda, 1957) and the region's flood

plains have suffered many rock avalanches. Even so, the area is populated by nearly 400,000 people. Understanding the region's geomorphic history is thus important for predicting future disasters.

The Ainoura plain is one of the region's small plains; several ruins have been excavated on this plain, indicating that people have been living on the Ainoura plain for at least 6 ka. The Monzen ruins represent one such site. The Nagasaki Prefectural Board of Education excavated the ruins

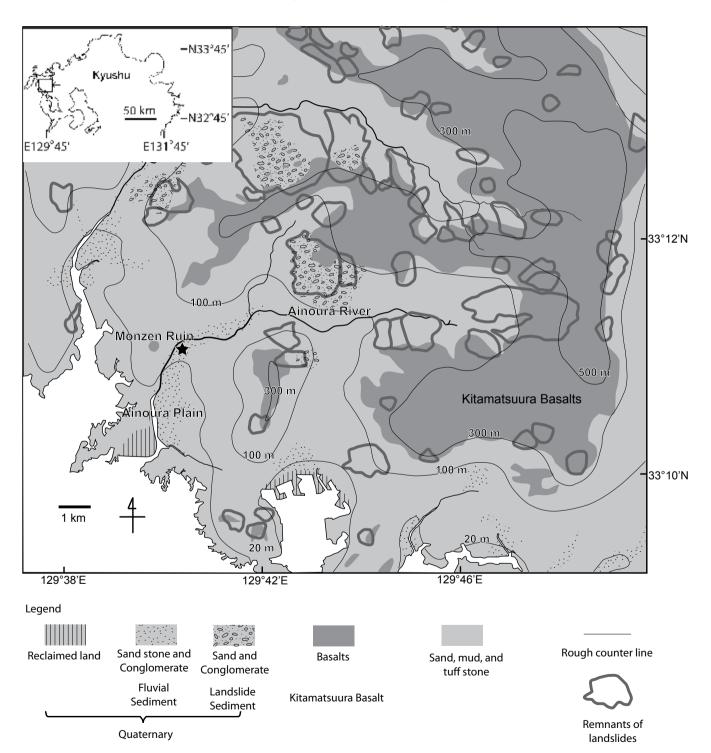


Figure 1. Topomap of the Ainoura Plain, Sasebo, and its location in Northern Kyushu, Japan. A detailed map around the Monzen Ruin is shown in Figure 2. Note that the higher portion of the mountain range is covered by basalts, and landslides have occurred around the boundary between the basalt layer and underlying Sand, mud, and tuff layer. Small islands other than Kyushu island are not depicted. Modified from Hatano et al. (1974) and Matusi et al. (1989).

and published a report (2006); this paper is partly based on their excavation work. Their results were used to reveal the plain's geomorphic history, including being struck by many rock avalanches.

Geological data

The complicated shoreline of north-western Kyushu, Japan, has resulted from active subduction (Shimoyama *et al.*, 1999) combined with erosion. This area lacks flatlands, but there are occasional wide river valleys; the Ainoura plain in western Sasebo, in which Monzen ruins are located, is one of these.

Figure 1 gives a geological map of the Ainoura plain and vicinity. The most widespread stratigraphic units in the region are the Sasebo group and the Ainoura group consisting of Paleogene sand and mud (Matui *et al.*, 1989). These groups lie beneath the 6-9 Ma "Kitamatsuura Basalt" (which includes basalt and andesite; Kaneoka and Aramaki, 1971; Matsui and Uto, 1998). Basalt and andesite cover mountain tops and the structure is unstable. Erosion of fragile sand and mud layers causes landslides; accordingly, this type of landslide has been named a Hokusho-type landslide by Noda (1957) and a detailed landslide map was prepared by Hatano *et al.* (1974a,b).

The Ainoura plain is an alluvial plain on the Ainoura River; the Hino River is a tributary of the Ainoura River and these two rivers converge near Ainoura Bay. Several ruins are present on the Ainoura plain (Figure 2). Aso (1972) reported discoveries of houses, shell and animal bone middens and tombs from the early Jomon era (c.a. 4,000 BC). 400 meters upstream of the Shimomotoyama cave ruin, the Shitanda ruins were built during the final Jomon to middle Yayoi era (c.a. 3,000 to 1,000 BC). Sasebo Municipal Council (SMC) excavated houses, paddy fields and stone tombs in this ruin from 1990 to 1995 (Kumura, 1994). Takebe C ruins were built on a hill and were excavated by the Nagasaki Prefectural Board of Education (NPBE) (Nagasaki Prefectural Board of Education, 2008). This included the Monzen ruins, located in the middle of the Ainoura plain (N 33°12'06", E129°40'09", WGS84), where the Ainoura River curves to the south (Figure 2). NPBE excavated the ruins and published a report in 2006 and 2008; they discovered archaeological remains from the initial Jomon era (c.a. 5,000 BC) to the modern era, such as earthen vessels, shell middens, houses and tombs. The wooden tombs from the Yayoi era (c.a. 300 BC) have special archeologically importance as they are some of the oldest wooden tombs found in Japan. Their excavation also unearthed clay pots, ceramic wares made in China and in the Kinki region (near Kyoto), as well as agricultural tools, such as a wooden spade. They also found at least five ancient channels from inside the Monzen ruins and unearthed several archaeological items from these channels' conglomerate deposits. Most items from the ruins showed little evidence of abrasion. NPBE and SMC have conducted research on plant residues and phytoliths from the Monzen and Shitanda ruins. Research on the Monzen ruins indicates that the site was a moist environment surrounded by evergreen broadleaf forest. However, the amount of rice phytolith was not enough to affirm rice cultivation (NPBE, unpublished). On the contrary, research on the Shitanda ruins revealed rice phytoliths, indicating that rice might have been cultivated there (Kumura, 1994).

The non-abraded items and existence of the rice phytolith suggested that people were living on the plain. However, these ruins are characterised by having been built on an alluvial plain which has experienced several destructive avalanches. The chronology of Japanese history, especially concerning when the Yayoi era began, is an active area of research, currently lacking consensus (Ozuka, 2000; Onishi, 2005; Yamamoto, 2007: Nakahashi and Iizuka, 2008). This research has been based on the conservative view that sets the start of Yayoi era around 2,300 years ago (Ozuka, 2000).

Methods

A geological cross-section of the Ainoura plain was first prepared along the Ainoura River and the Hino River from drill core data provided by SMC, Northern Nagasaki Prefecture Development and Promotion Bureau and the Fujinaga Geological and Construction Corporation. A topographic cross-section was created based on a 50 m-mesh digital elevation model published by the Geological Information Authority of Japan (2002), using Kashmir 3D software.

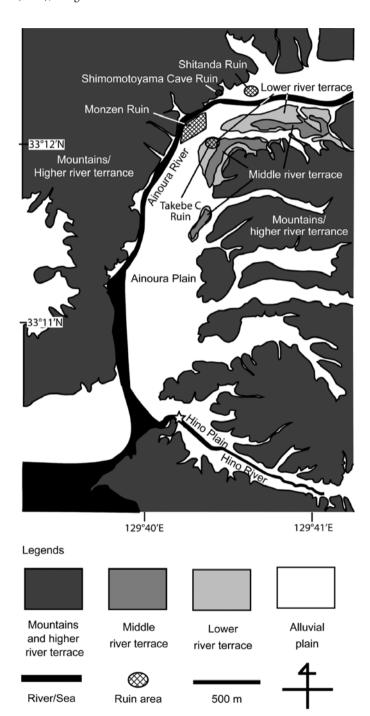


Figure 2. Terrain classification in the Ainoura plain showing the major rivers. The Monzen ruins occupy a large portion of the narrow part of the floodplain surrounded by mountains and river terraces. Mollusk shells were collected from the location indicated by the stars.

The paleoshoreline's approximate position was determined by using drill data: Holocene deposits including marine shells and others which were continental. The location of the paleoshoreline was estimated from the intersection of the lowest marine deposits with the current terrain. The Holocene marine layer's maximum extension represents the paleoshoreline during the Holocene climatic optimum (4000BC), when the sea level was at its highest. Shell fossils were collected from the bank of the Hino River

(Figure 2). The plain's depositional history was investigated by plotting sediment layers' elevation changes. The ages of layers determined by NPBE (2008) was used mainly by means of seriation of potteries. When cultural objects were not found, ¹⁴C from fallen trees and tephra derived from the Sakurajia volcano was also used. Referring tephra helped to solve the problem of secondary deposition in some players. The elevations were adjusted by applying the modern gradient of the Ainoura River near the ruins,

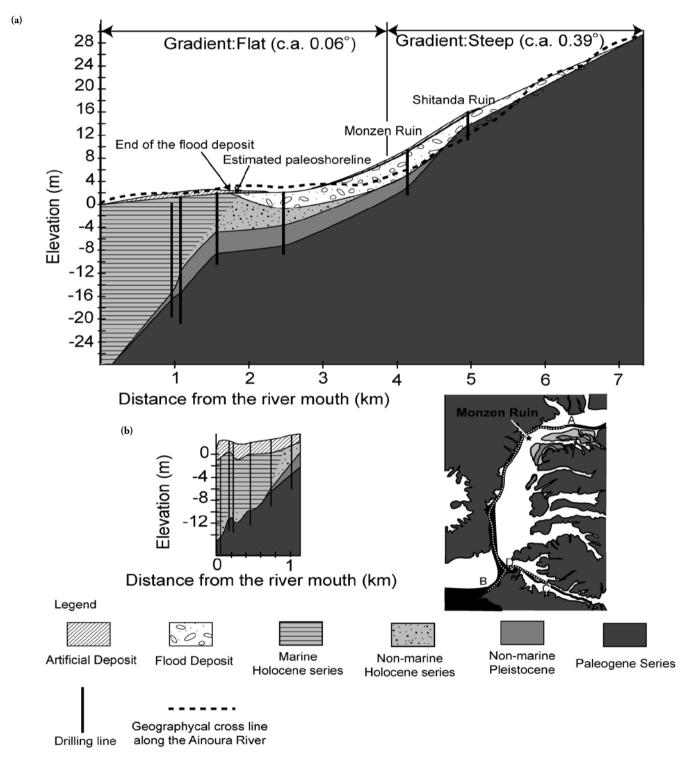


Figure 3. Geological and topographic sections along the Ainoura (a) and Hino rivers (b). Drilling logs were used to draw the closed circles and triangles to draw a geological cross-section of Figure 4; the former is along the transect AB and the latter is along the transect CD of the small topographic map. Flood deposits in the Ainoura plain could not be found in the Hino River flood plain. The Monzen ruins are located near the sudden reduction in the gradient.

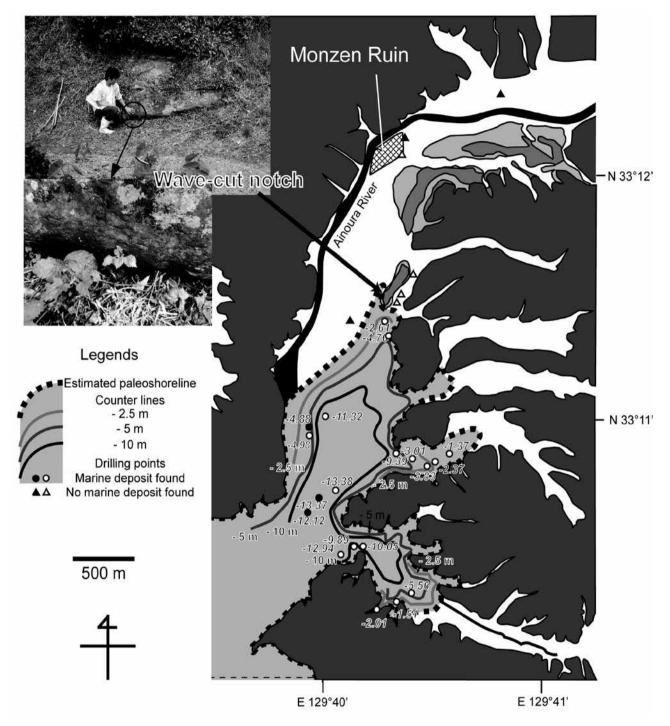


Figure 4. The distribution of the drilling data and the estimated paleoshoreline of the Ainoura plain. The paleoshoreline was estimated based on marine deposit distribution, shown as the intersection of the lowest marine deposits with the current terrain. The marine deposit was reported from locations marked with closed and open circles, but not from locations marked with closed and open triangles. The number in italics is the marine layer elevation at each location.

measured as 3.89/1,000 using an auto level. Accordingly, 38.9 cm would be added to the elevation of a sediment layer 100 m downwards from the centre of the ruins along the paleoriver channel, for example.

Results

A cross-section (Figure 3) of the Ainoura plain showed that the Monzen ruins were located near a sudden reduction in the gradient and that the flood deposit layer was relatively thick here. Figure 3A also represents the cross-section of the Ainoura plain. From bottom to top, the plain

consists of a Paleocene unit, a non-marine Pleistocene unit, a non-marine Holocene unit, flood deposits after the Holocene climatic optimum and human-made deposits. Many basaltic conglomerates were included in this plain in the non-marine Holocene unit. A geological cross-section of the Hino River is shown in Figure 3B. Although layers similar to those of the Ainoura River were found in this region, no flood deposits were observed. The following mollusk shells were collected from the bank of the Hino River: granosa, Macoma tokyoensis Makiyama, Anadotia stearnsina, Pephia undulata, Pistris capsoides, Batillaria zonalis, Cerithideopsilla cingulata, Turbo cornatus coreensis, Cryptonatica tigrina, Cerithideopsilla djadjariensis,

Nassarius livescens and Hemifusus ternatanus. Most of these mollusks lived in the intertidal zone.

The 4,000 BC paleoshoreline was estimated (Figure 4). Beside the drill data, a wave-cut notch found on a hillside was considered as evidence of the paleoshoreline. Figure 4 shows that the Monzen ruins were about 1 km away from the shoreline; the shoreline at the Hino River was about 1 km away from the current shoreline.

Figure 5 shows how mud, sand and conglomerate layers developed on the plain. Figure 5A show that sand and mud layer levels from the initial Jomon era were 1 to 2 m lower than those of later periods. The levels of sand and mud layers rose suddenly about 4,000 BC, then stopped rising. Figure 5B shows that levels of conglomerate layers have risen since the last 5,000 BC and four to five debris flows presumably brought the gravels striking the Monzen ruins.

Discussion

The Ainoura plain's Paleogene unit consists of sand and mud layers having nearly horizontal stratification, covered by Kitamatsuura basalts. Landslides easily occur in such an area when underground water pressure rises. Noda (1957) reported that the area has experienced many such landslides. Comparing the cross-sections of the Hino and Ainoura rivers (Figure 3) revealed differences in their depositional histories. Considering the shell fossils from the Hino River, it was inferred that the Hino plain was originally an intertidal flat dominated by marginal marine sedimentation that subsequently became damp, subaerially exposed ground. By contrast, basaltic gravels from the Ainoura plain must have either been derived from landslides near the ruins, or carried from farther upstream. Since there is no evidence of landslides near the ruins (Figure 1), it was assumed that the basaltic gravels of the Monzen ruins were carried by the Ainoura River; fluvial influence is thus more pronounced than marine influence on the Ainoura plain.

A geological cross-section of the Ainoura Plain (Figure 3) revealed that the Monzen ruins were located near the sudden reduction in the gradient, and where the flood layer is thick. This suggested that such location was where the river's flow rate suddenly drops and sedimentation occurs. These gravels did not begin to be deposited until 1,000 BC (Figure 5B) at the location of the Monzen ruins. Gravel became successively deposited over the conglomerate layers after that, leading to the rising of the riverbed. Such trend was dissimilar to that of mud and sand layers (Figure 5A). Mud and sand layers most likely responded to changes in river water level which was influenced by changes in sea level. Yokoyama et al., (1999) reported sea level change in Tsumizu, western Kyushu, 50 km south-east of the Ainoura plain. The sea level change there was characterised by a sudden rise in sea level around 5,000 to 3,000 BC and a subsequent gradual fall, similar to the sand-mud deposition pattern on the Ainoura plain. Since the geomorphic history of western Kyushu is complicated (Yokoyama et al., 1996; Shomoyama et al., 1999), one cannot be certain whether the Ainoura plain experienced the same change in sea level as at Tsumizu, but it is possible that changes in Ainoura plain elevation record/reflect changes in sea level in western Kyushu.

Figure 4 shows that the paleoshoreline was located about 1-2 km seaward from the Monzen ruins. After the Holocene climatic optimum, the shoreline has been gradually prograding due to sedimentation; however, the cross-section (Figure 3) showing locations between the paleoshoreline and the end of flood deposition is quite close. It was thus concluded that progradation did not alter the shoreline much, at least until the 17th century, when the Japanese began to reclaim the land (Nagata 1994). It can thus be expected that the deposition which happened at the location of the Monzen ruins described above probably also happened in the region between the Monzen ruins and the estimated paleoshoreline, accounting for almost half of the Ainoura Plain.

Conclusions

This research was focused on the geomorphic history of the Ainoura plain based on the Monzen ruins. Most of the strata of the Ainoura plain consists of basaltic conglomerate (flood deposit) and sand. Such basaltic gravels were derived from the Kitamatsuura basalts in the hinterland; such gravel represents landslide debris carried by floods to the plain to be deposited. The geographical cross-section revealed that the Monzen ruins were located at the sudden reduction in the river gradient, more flood sediments thus being deposited there. It was also found that mud and sand first filled in the trough, probably as part of the Jomon transgression caused by the Holocene climatic optimum. Later, conglomerate began to replace sand and mud layers by successive erosion and sedimentation, thus raising the riverbed. Around the final Jomon era (c.a. 500 BC), depositional elevation levels of gravel surpassed those of sand and mud. Considering the position of the paleoshoreline, such scenario is probably applicable to about half of the Ainoura plain.

Acknowledgments

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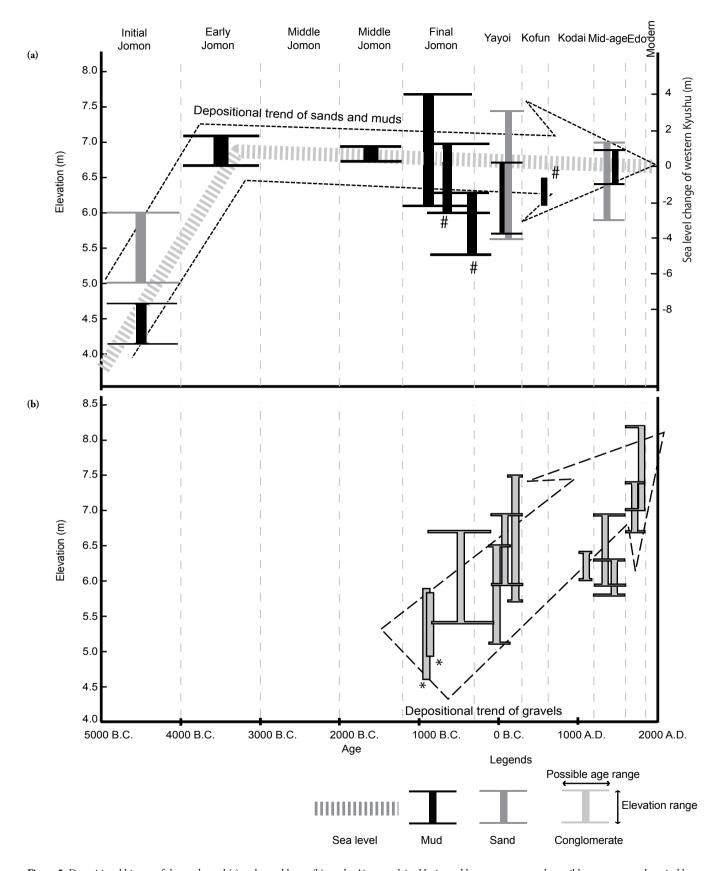


Figure 5. Depositional history of the sand, mud (a) and gravel layers (b) on the Ainoura plain. Horizontal bars represent rough possible age ranges and vertical bars represent altitudes for layers connected along the river gradient. Cross-section diagrams prepared during excavations were referred to for the elevations of the layers. The layers' ages were determined by seriation, 14C dating and tephra. Bars marked with an * are layers whose ages were determined by 14C and # by tephra. Sand and mud layers' depositional elevations rose around 5,000 to 4,000 BC, thereby agreeing with the change in sea level. The gravel layers began to be deposited around 1,000 BC, then already becoming replaced by deposited sand and mud layers.

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